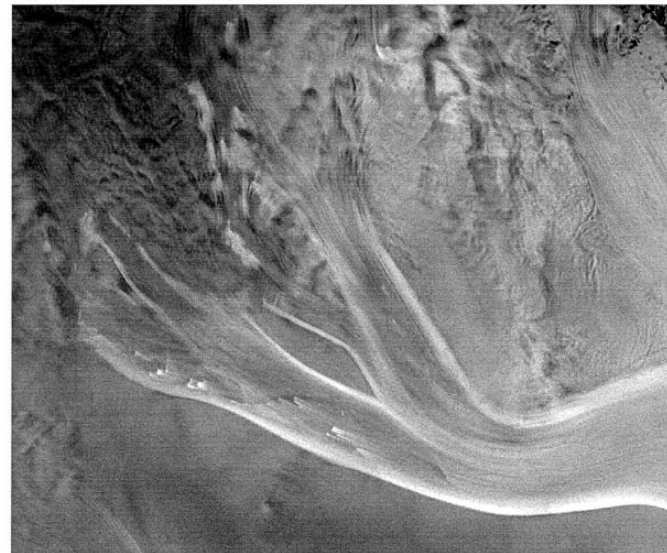


Deformation-induced melting in the margins of the West-Antarctic ice streams

Jenny Suckale
Harvard SEAS

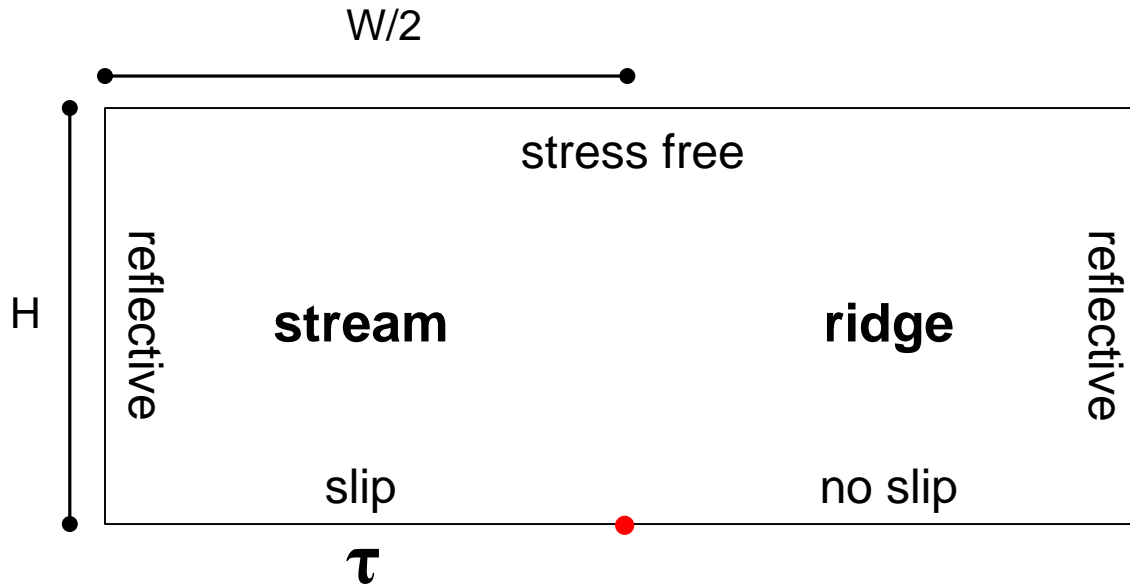
Collaborators:

John D. Platt, Harvard, SEAS
Thibaut Perol, Harvard, SEAS
Jim R. Rice, Harvard, SEAS



Scale
50 0 Kilometers

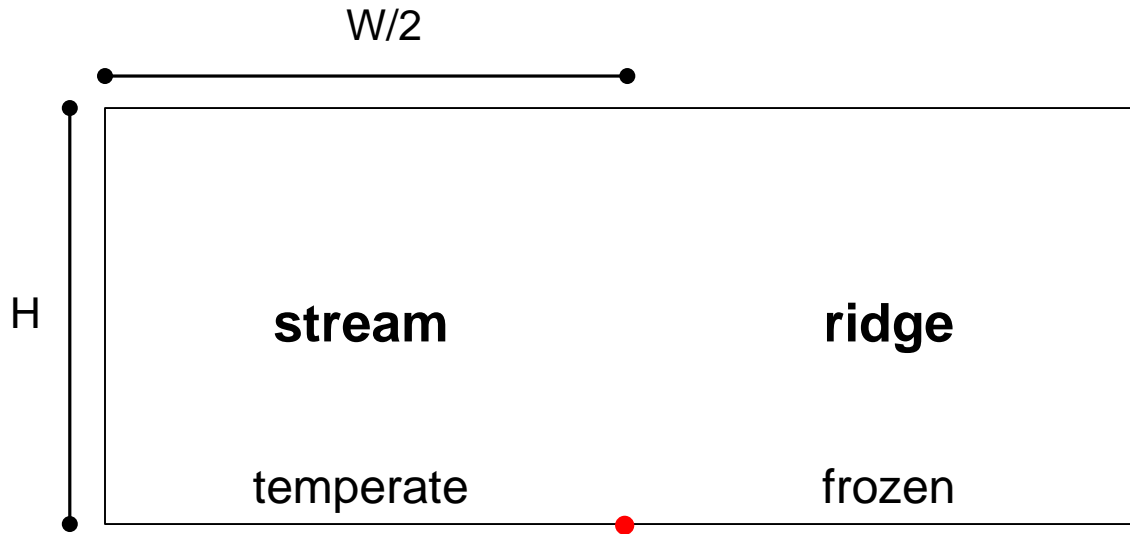
Mechanical model



$$\frac{\partial}{\partial y} \left(\mu \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) + \rho g \sin(\alpha) = 0$$

$$\dot{\epsilon} = \dot{\epsilon}_D + \dot{\epsilon}_G = A_D \tau + A_G \tau^3$$

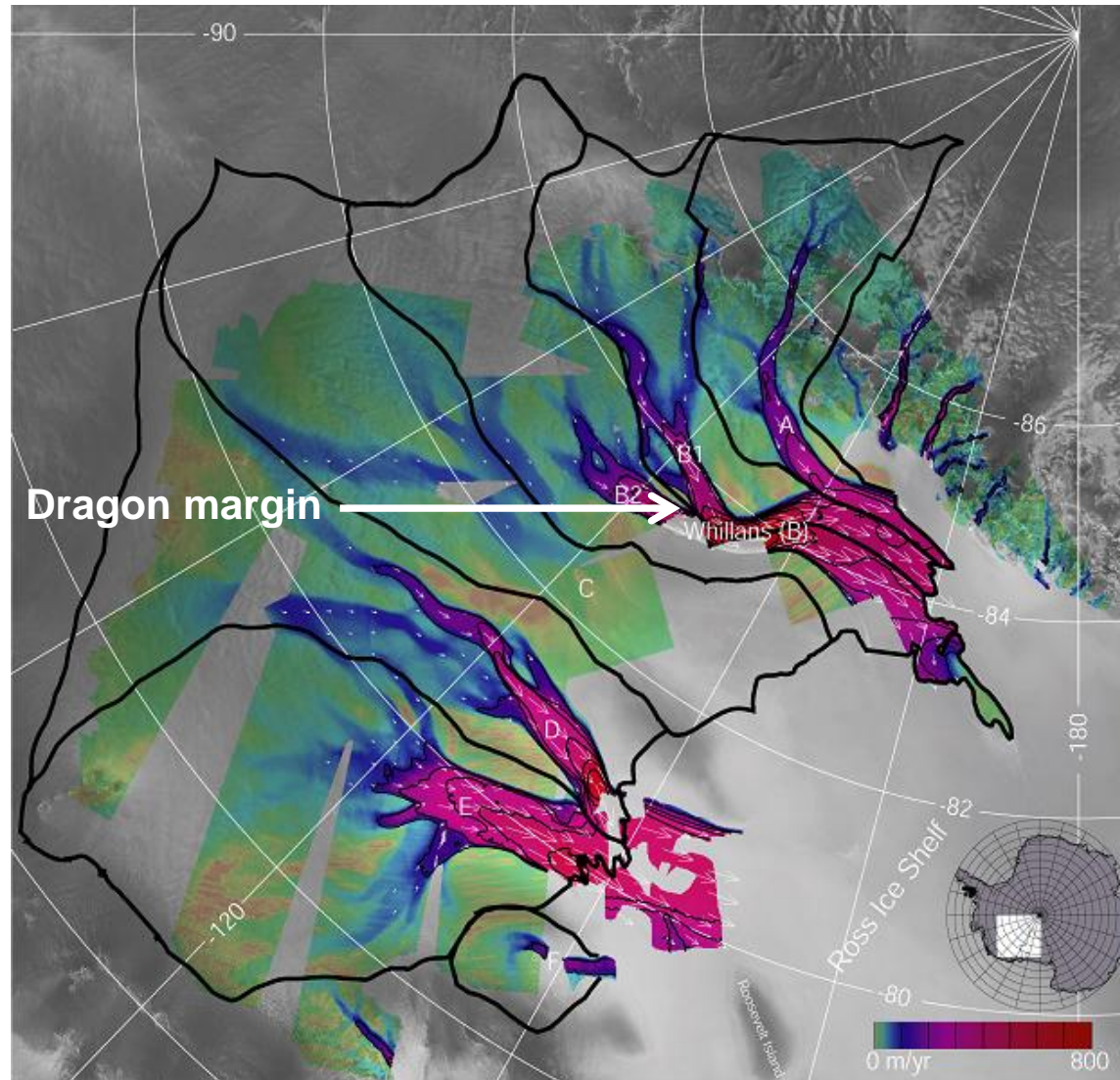
Thermal model



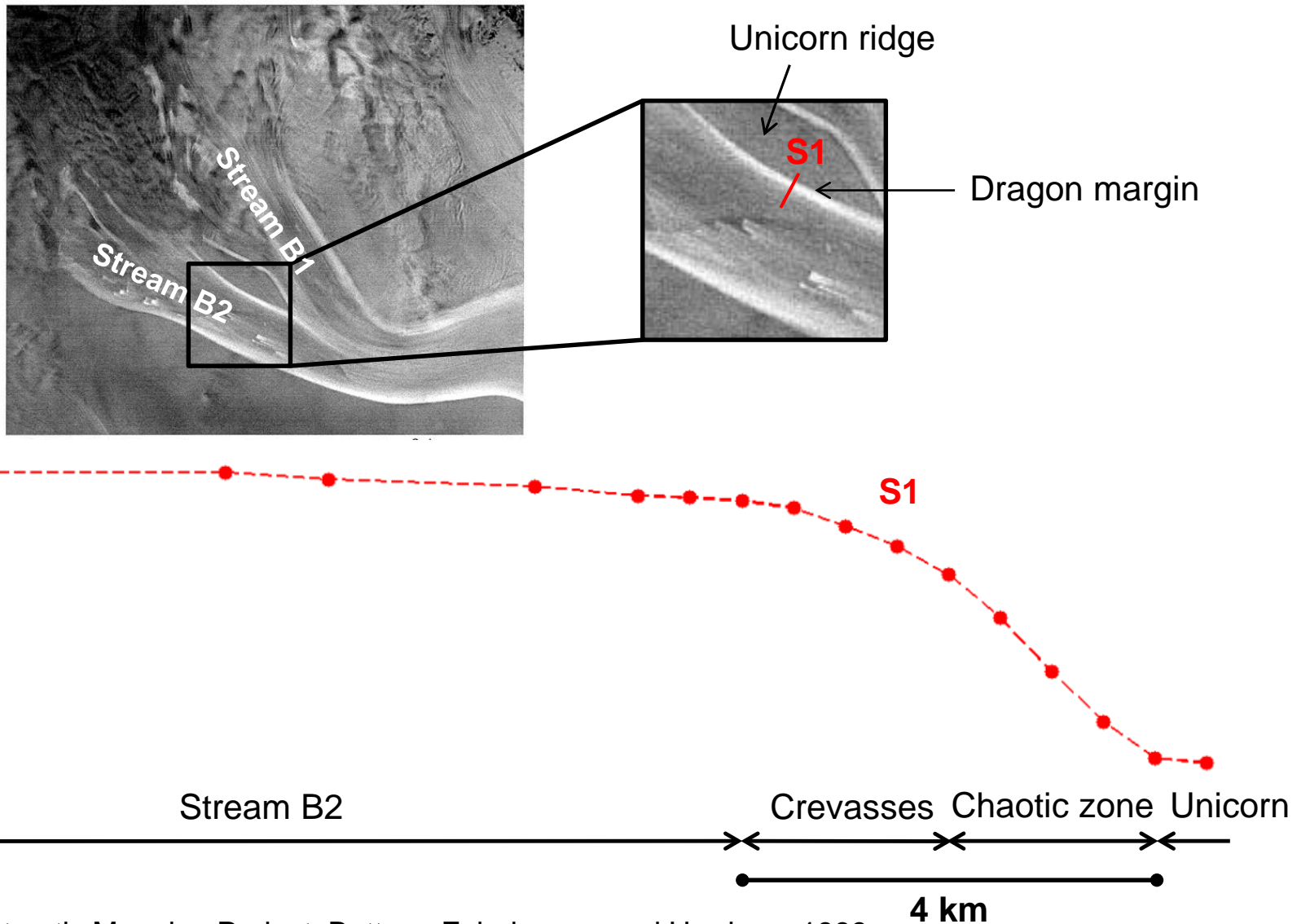
$$\frac{\partial}{\partial y} \left(k \frac{\partial T}{\partial y} \right) + \frac{\partial}{\partial z} \left(k \frac{\partial T}{\partial z} \right) + \rho c \left(v \frac{\partial T}{\partial y} + w \frac{\partial T}{\partial z} \right) + 2\tau_E \dot{\epsilon}_E - L\dot{m} = 0$$

$$\frac{\partial}{\partial y} \left(k \frac{\partial T}{\partial y} \right) + \frac{\partial}{\partial z} \left(k \frac{\partial T}{\partial z} \right) + \rho c \left(v \frac{\partial T}{\partial y} + w \frac{\partial T}{\partial z} \right) + (1 - H(T - T_m)) 2\tau_E \dot{\epsilon}_E = 0$$

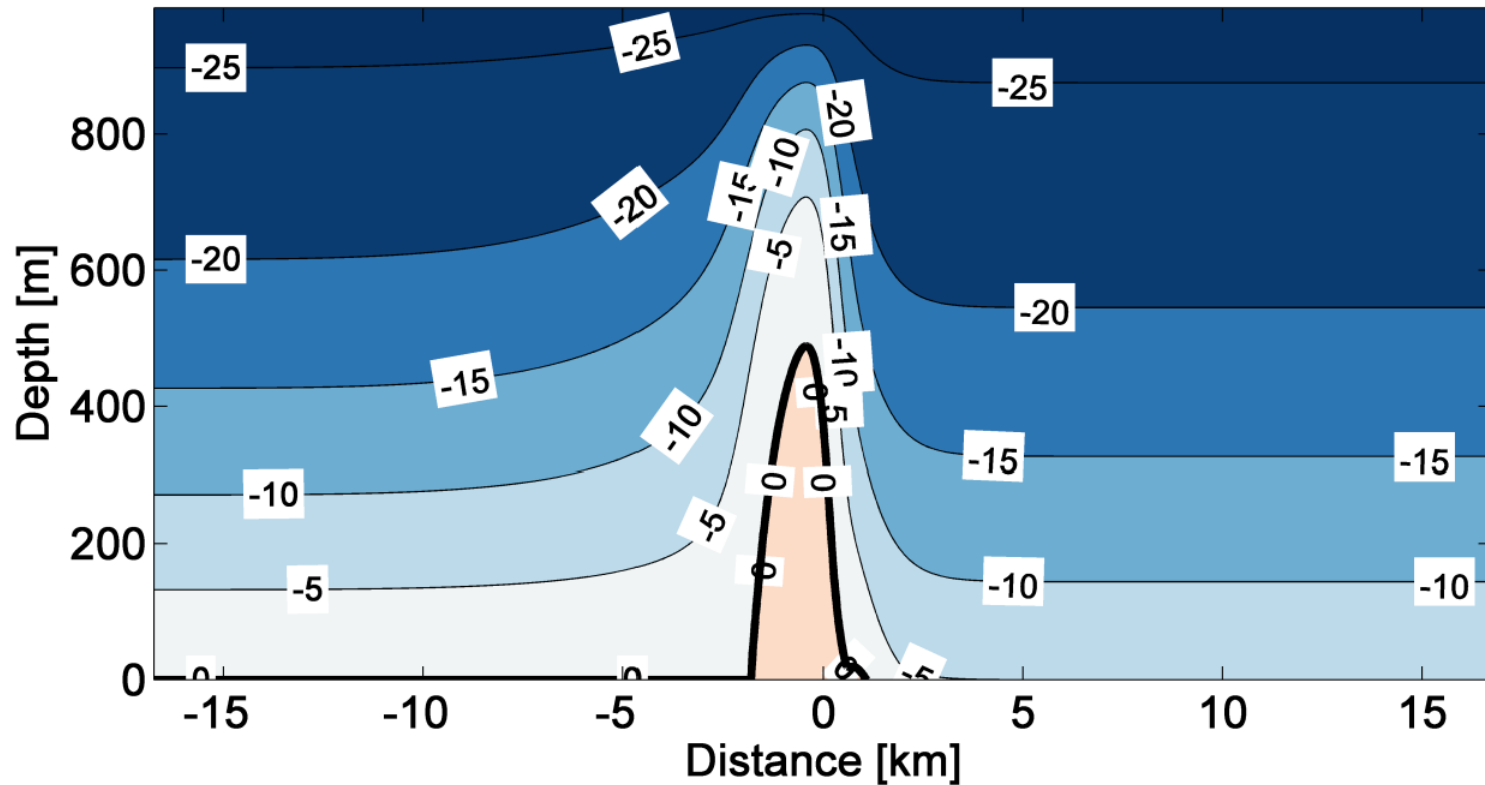
Test case: Dragon margin



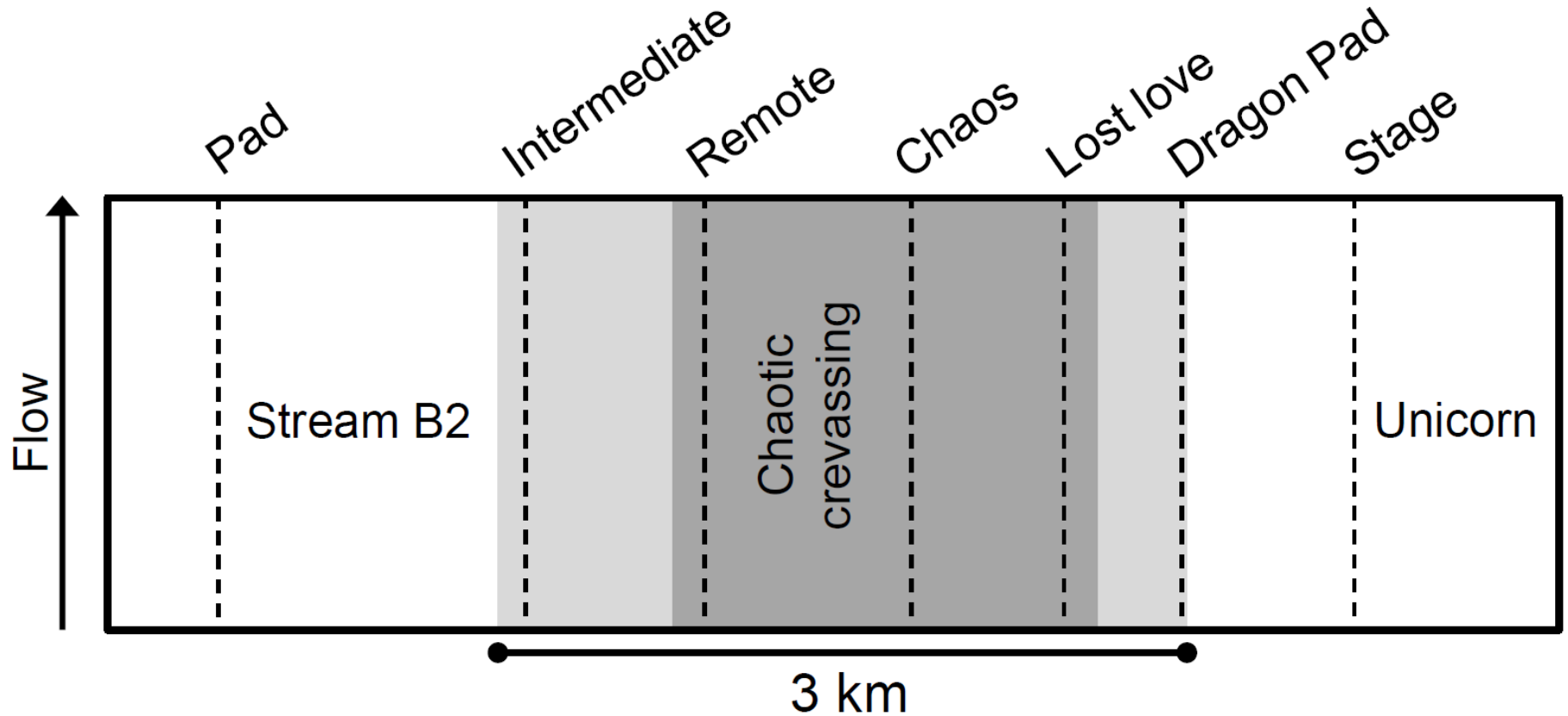
Testing the model – Velocity data



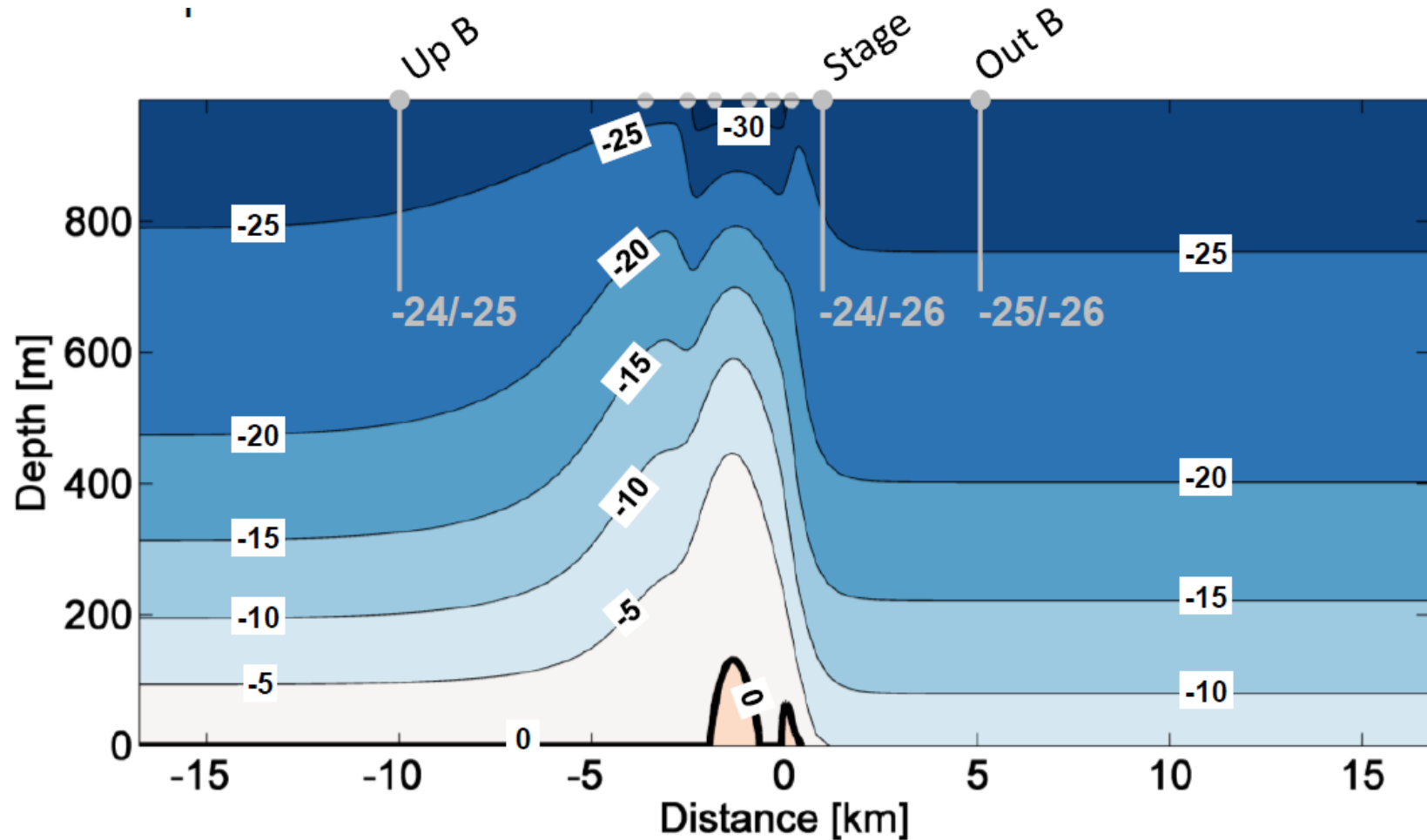
Reproducing velocity data



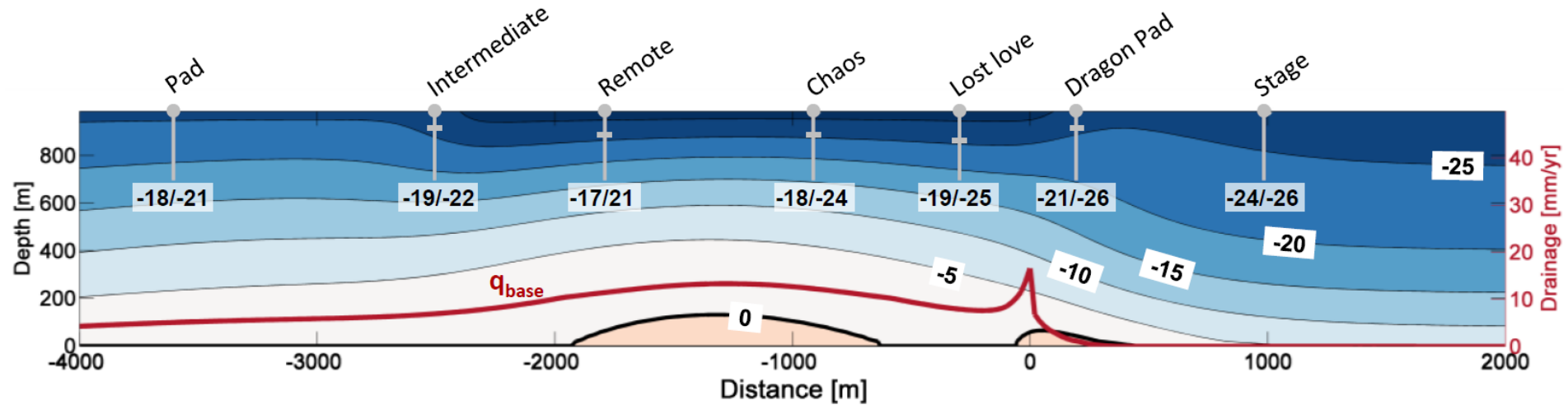
Testing the model - Temperature data



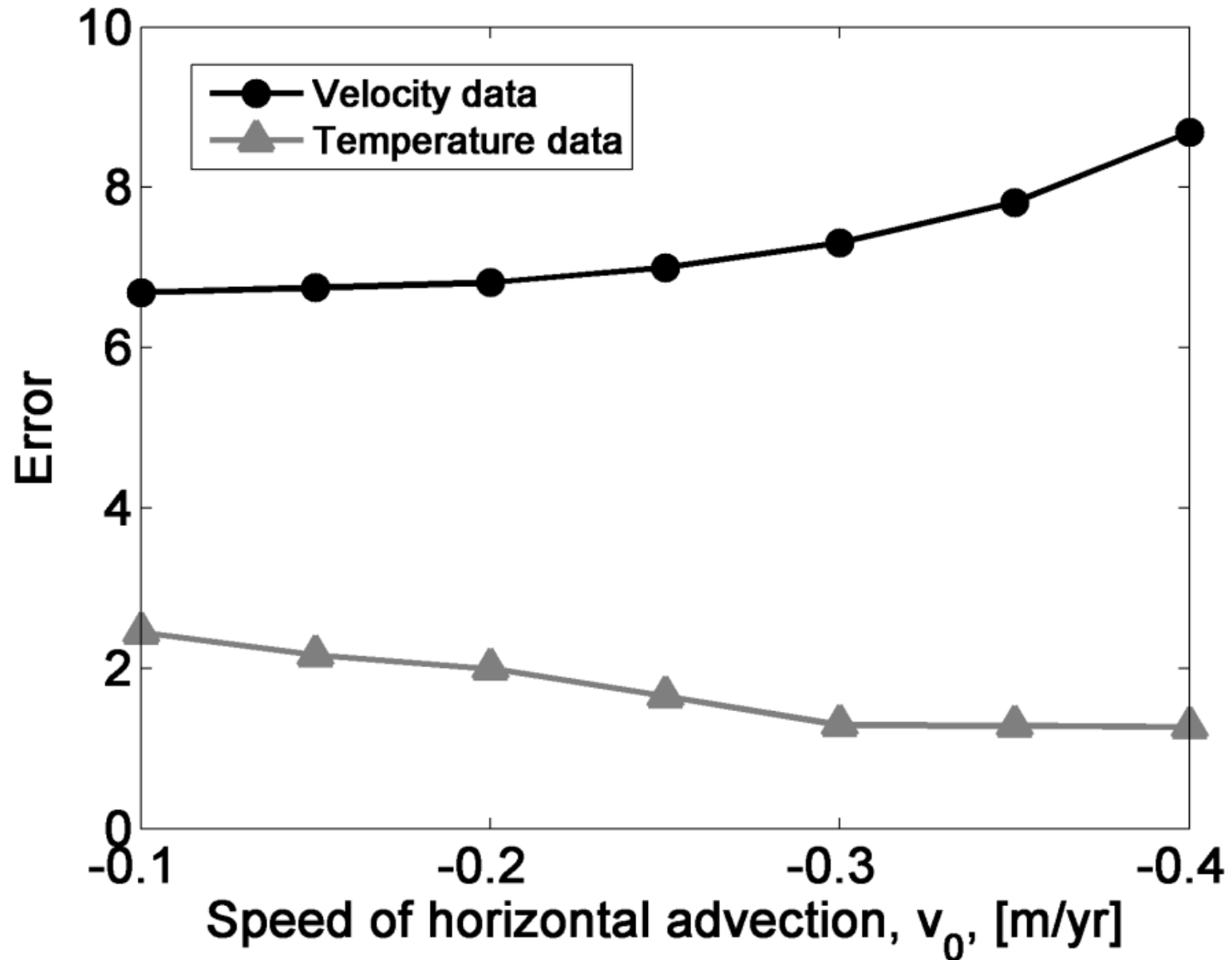
Reproducing temperature (far-field)



Reproducing temperature (near-field)

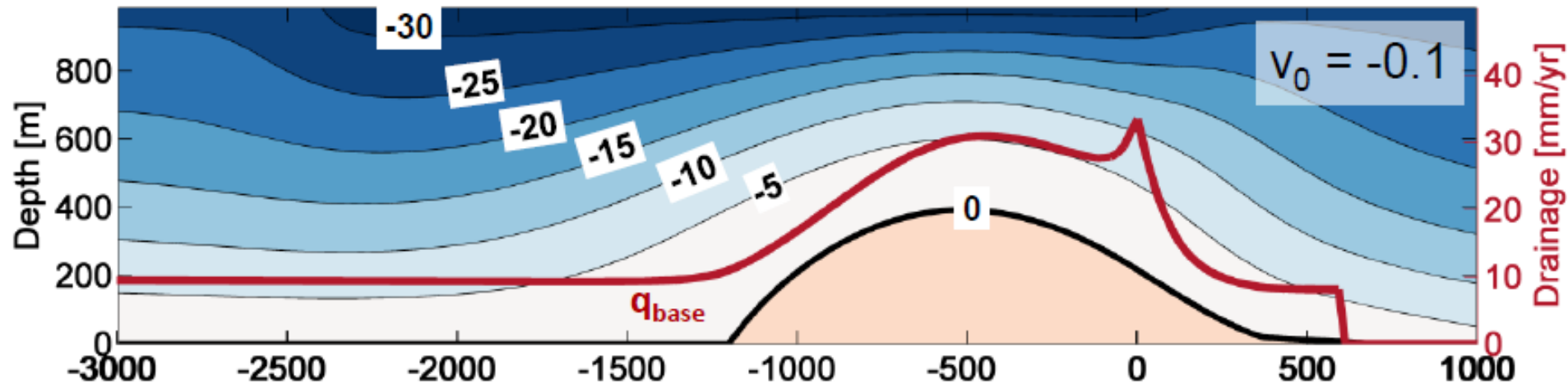


Error in reproducing both data sets simultaneously

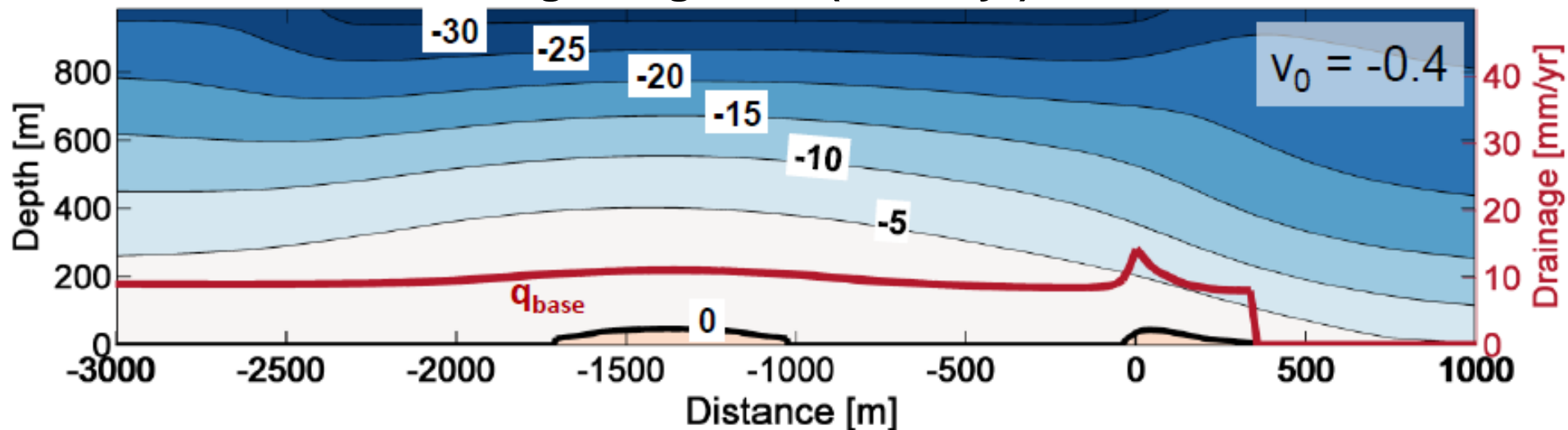


Temperate zones for different migration rates

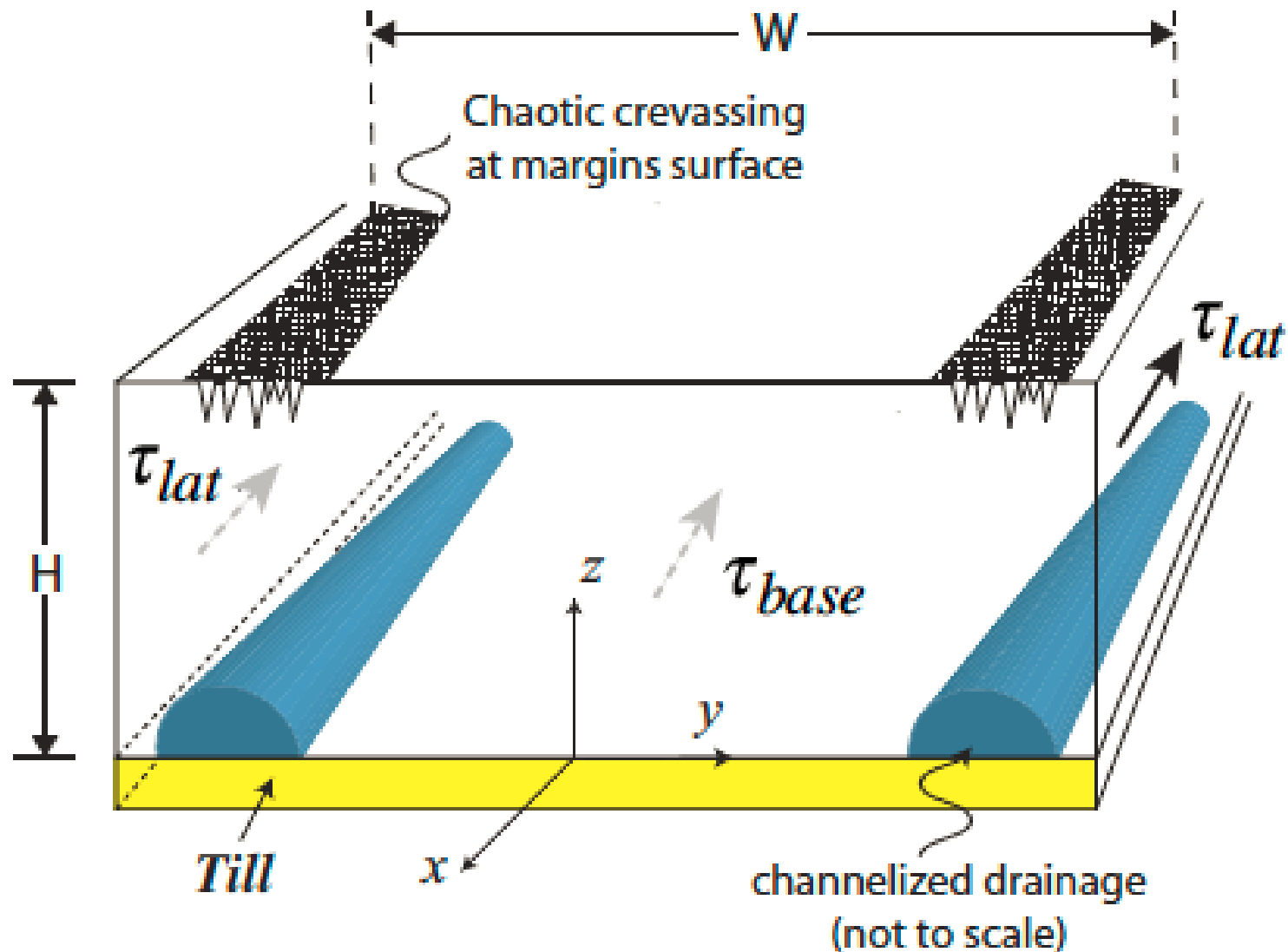
Slow rate of outward margin migration (-0.1 m/yr)



Fast rate of outward margin migration (-0.4 m/yr)

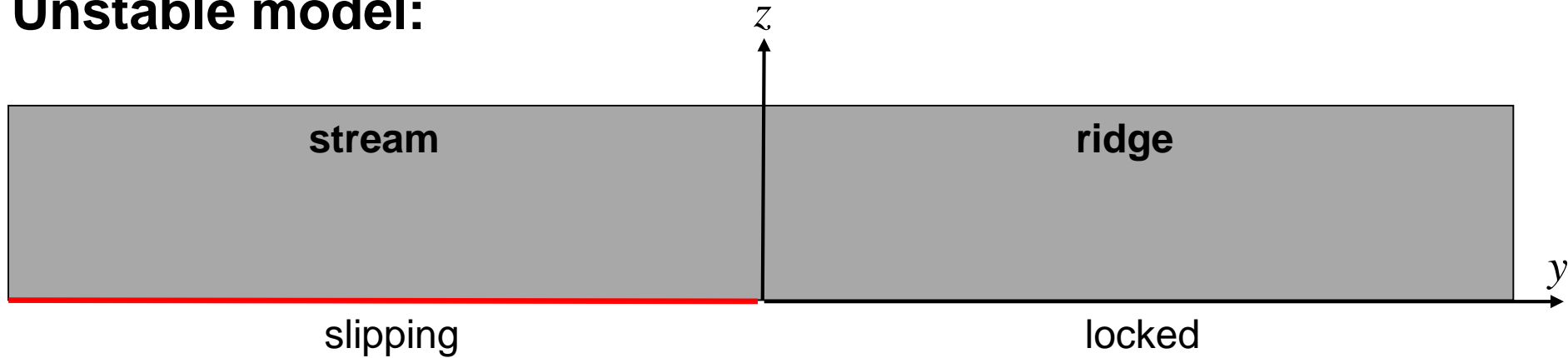


Channelized drainage systems in shear margins?



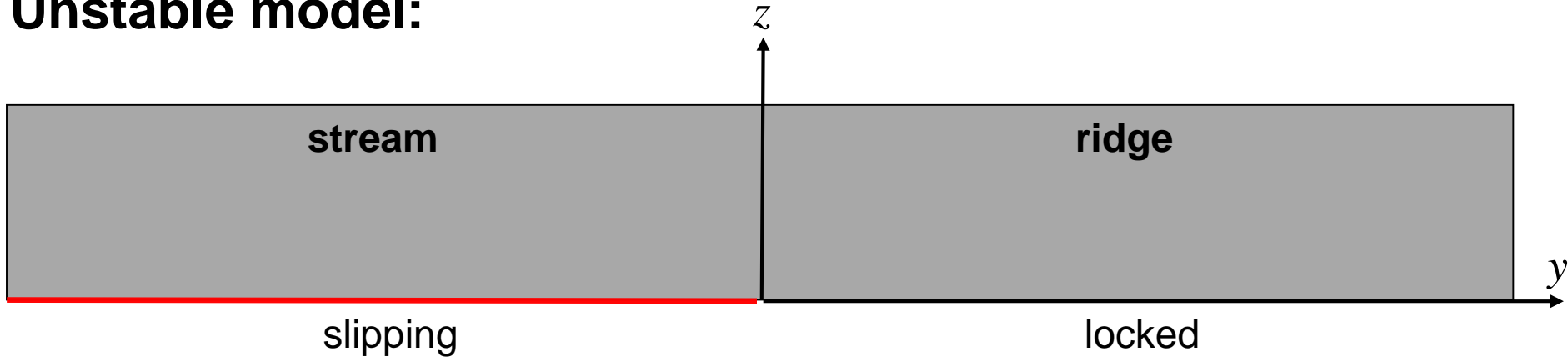
Dynamic ramifications of channel formation

Unstable model:

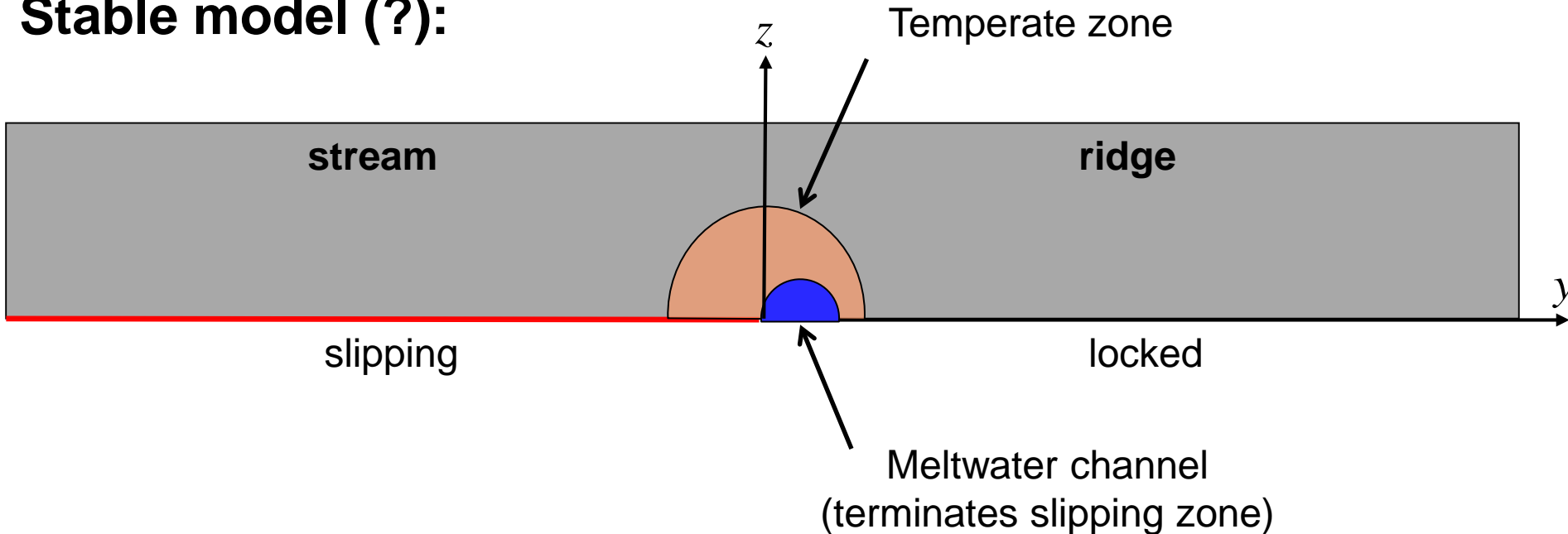


Dynamic ramifications of channel formation

Unstable model:



Stable model (?):

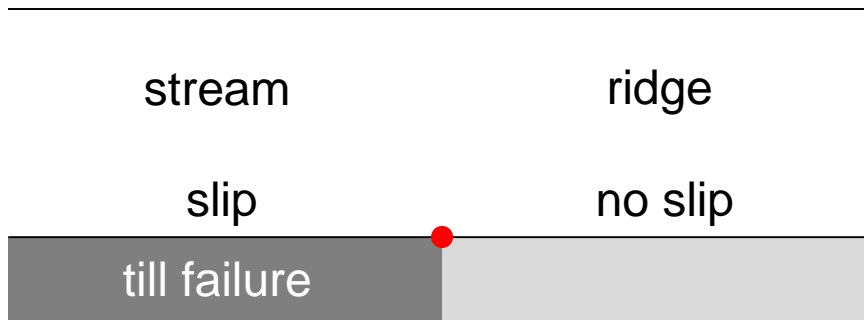


Conclusions

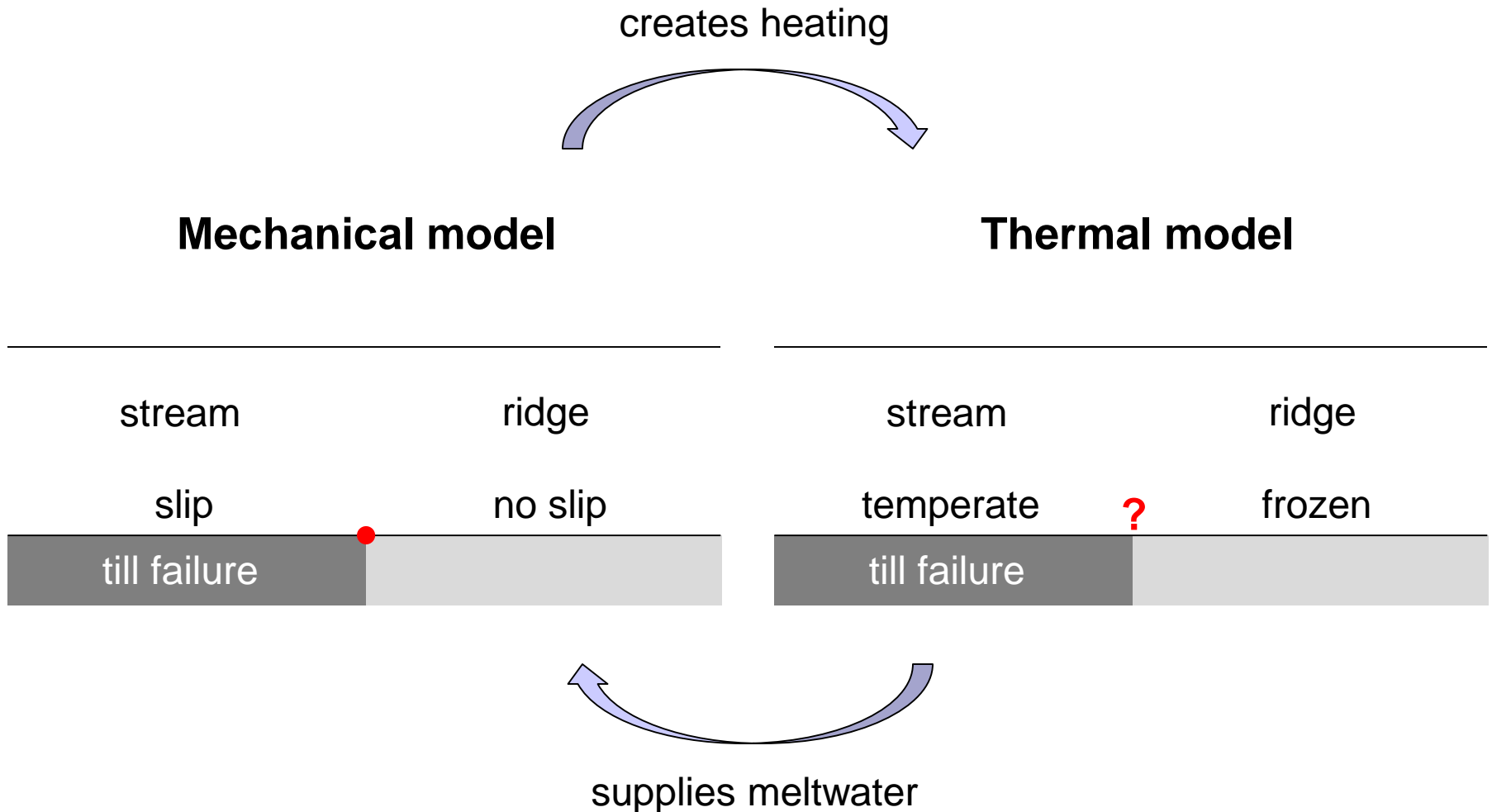
- 1. The shear margins of active ice streams are probably partially molten.**
- 2. We expect creation of significant amounts of meltwater.**
- 3. Where the meltwater is generated depends sensitively on the migration rate of the margin.**
- 4. Channel formation could explain locking.**

Thermomechanical coupling

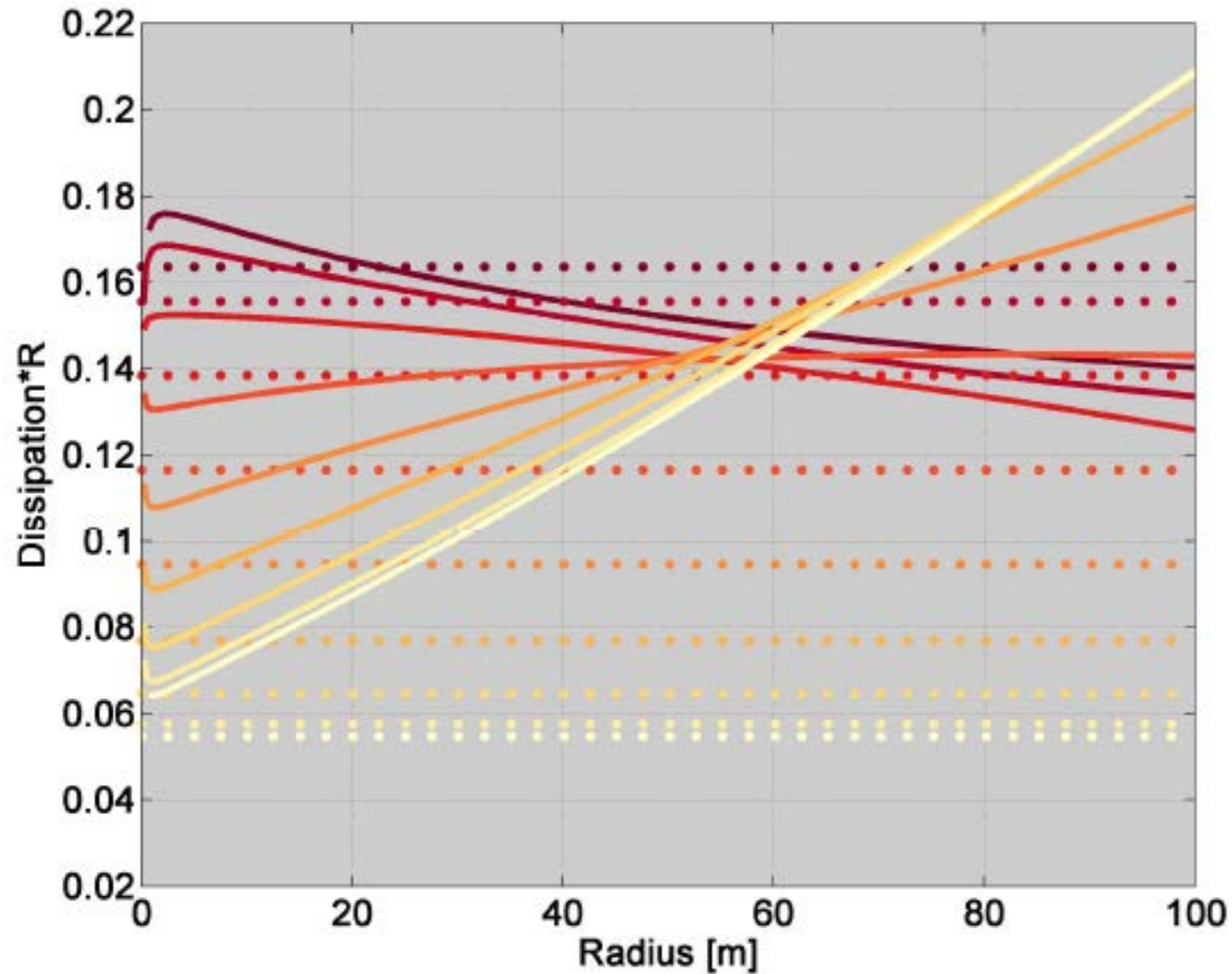
Mechanical model



Thermomechanical coupling

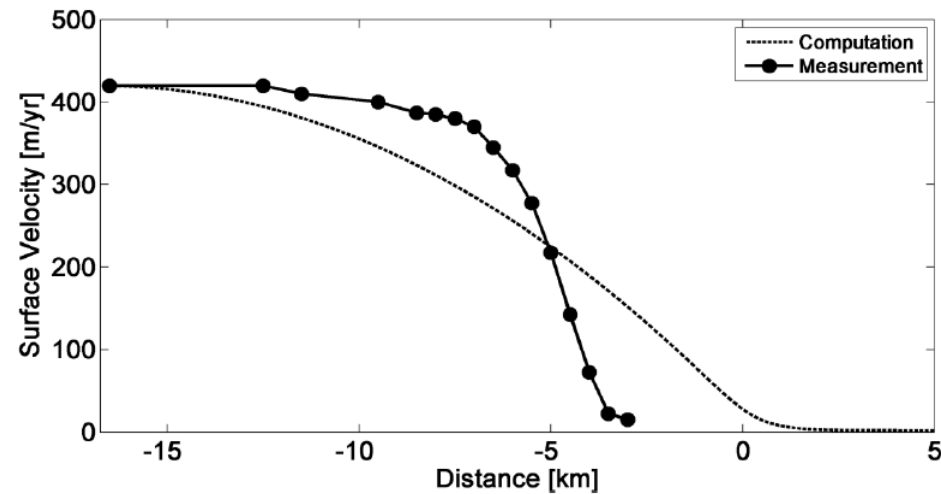
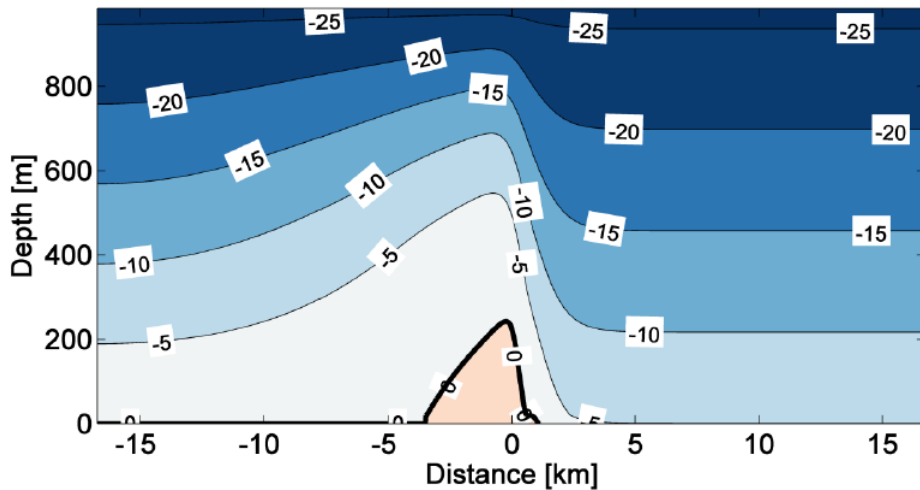


Benchmark computation

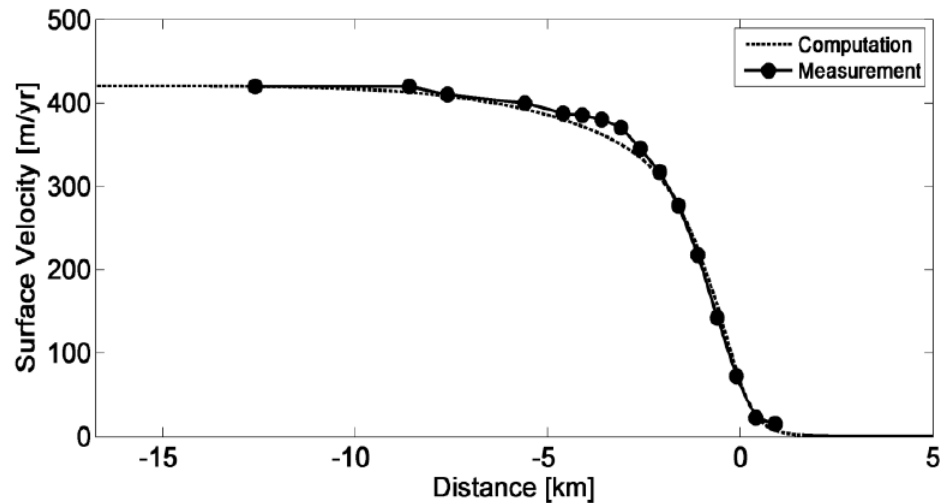
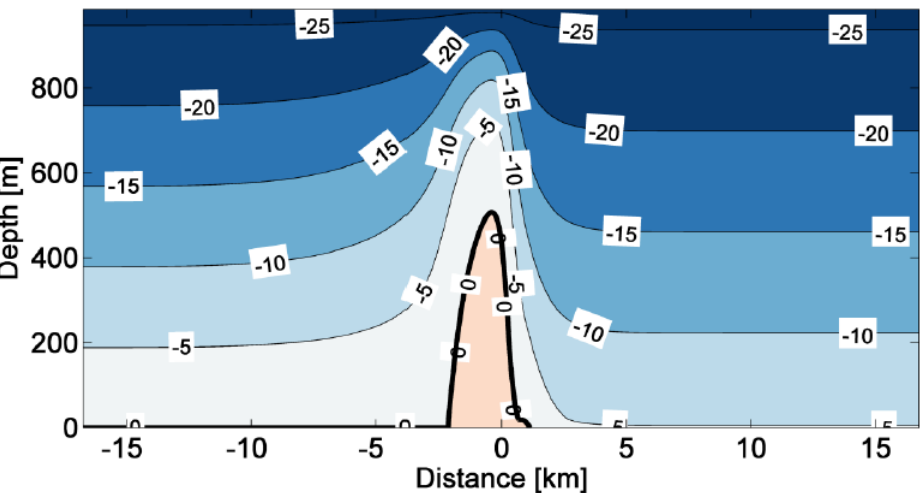


The importance of rheology

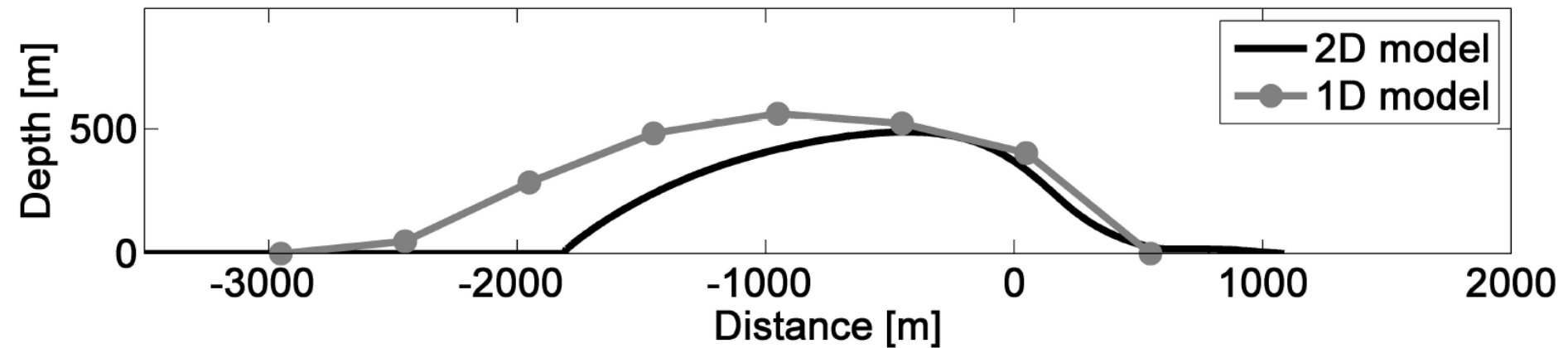
Newtonian rheology



Creep rheology



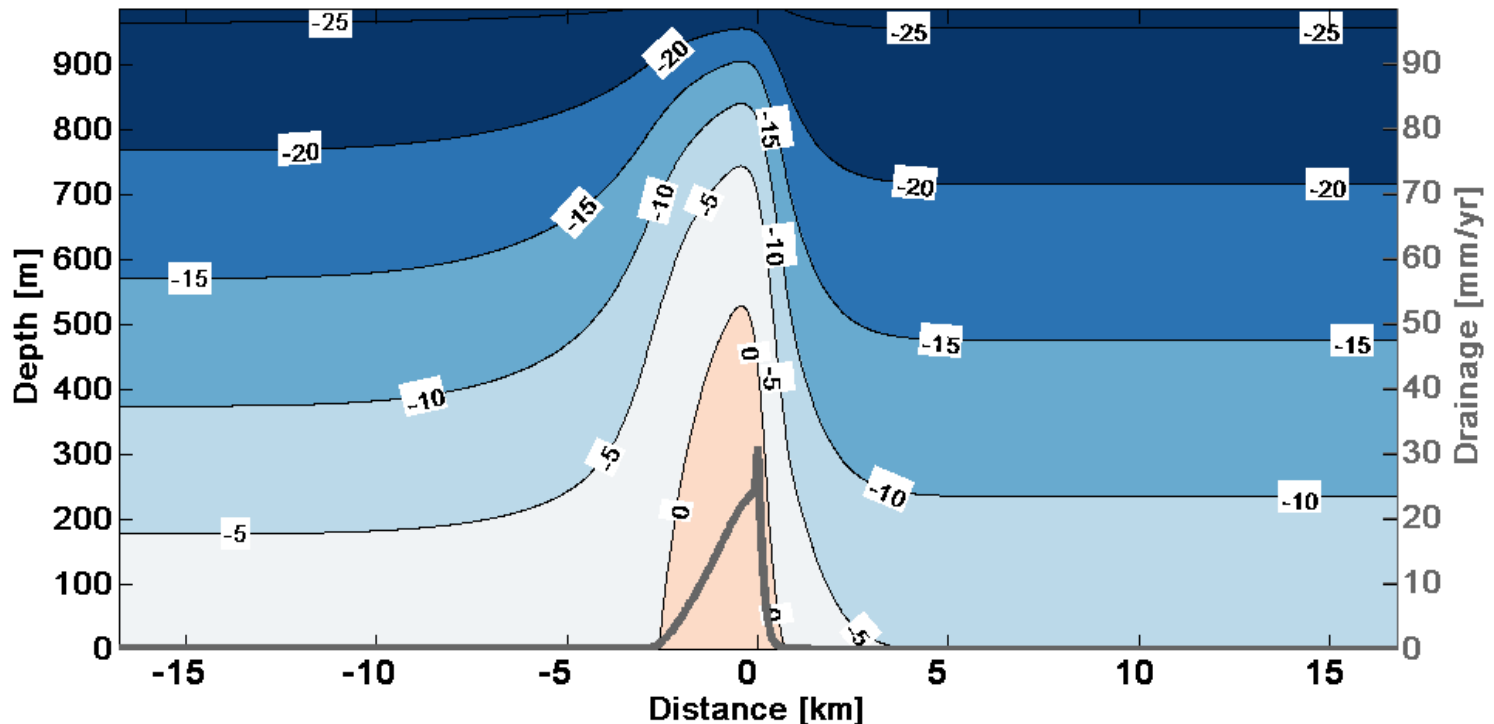
Comparison with 1D model (from Perol and Rice)



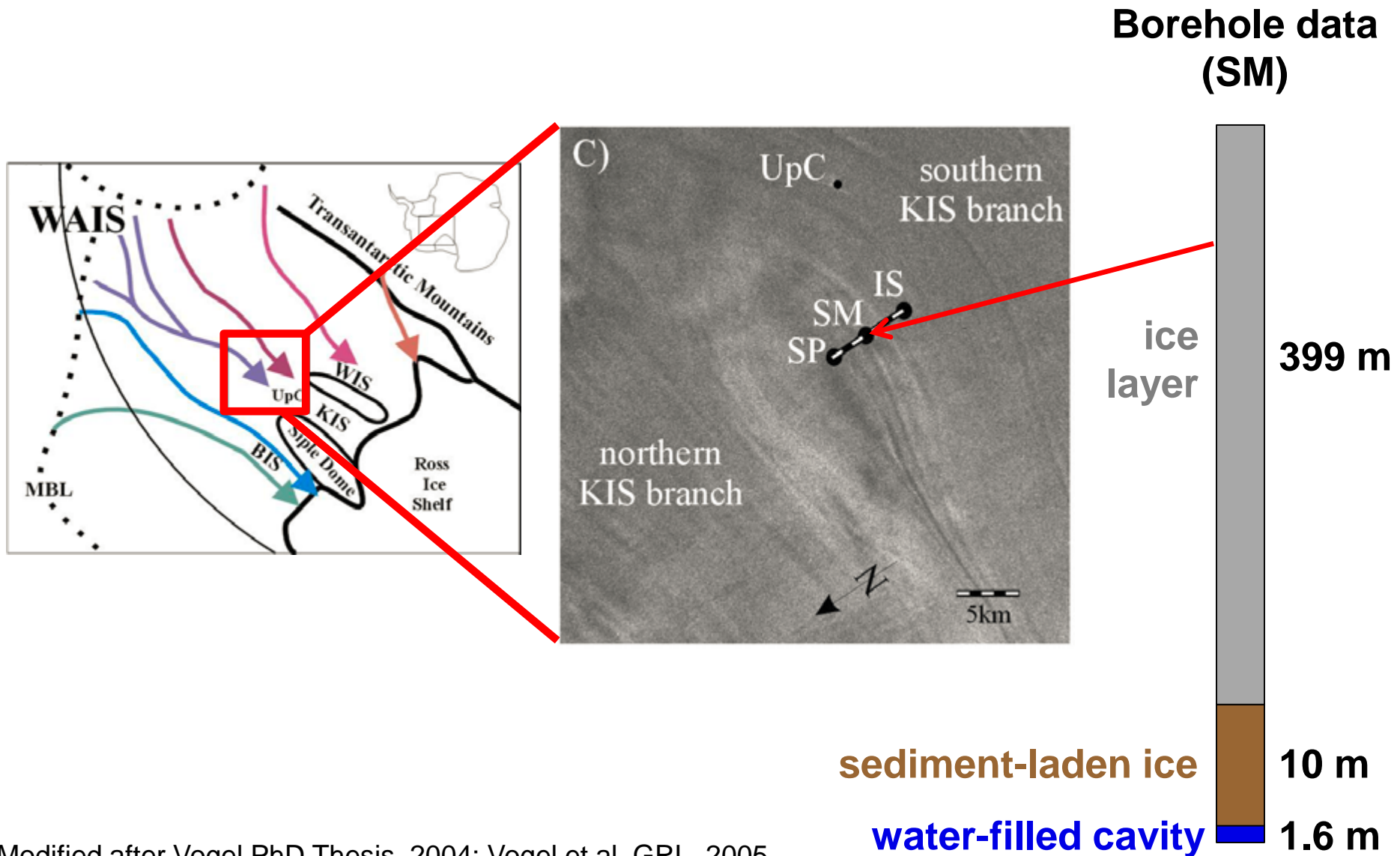
Estimating meltwater production

Mass balance:
$$\frac{\partial q_y}{\partial y} + \frac{\partial q_z}{\partial z} = -\frac{2\tau\dot{\epsilon}}{\rho_w L}$$

Meltwater flux:
$$q_{base} = -\int_0^{H_m} \frac{\dot{m}}{\rho_w} dz = -\int_0^{H_m} \frac{2\tau\dot{\epsilon}}{L\rho_w} dz$$



Observational evidence (Kamb ice stream)



Observational evidence (Clarke et al. 2000)

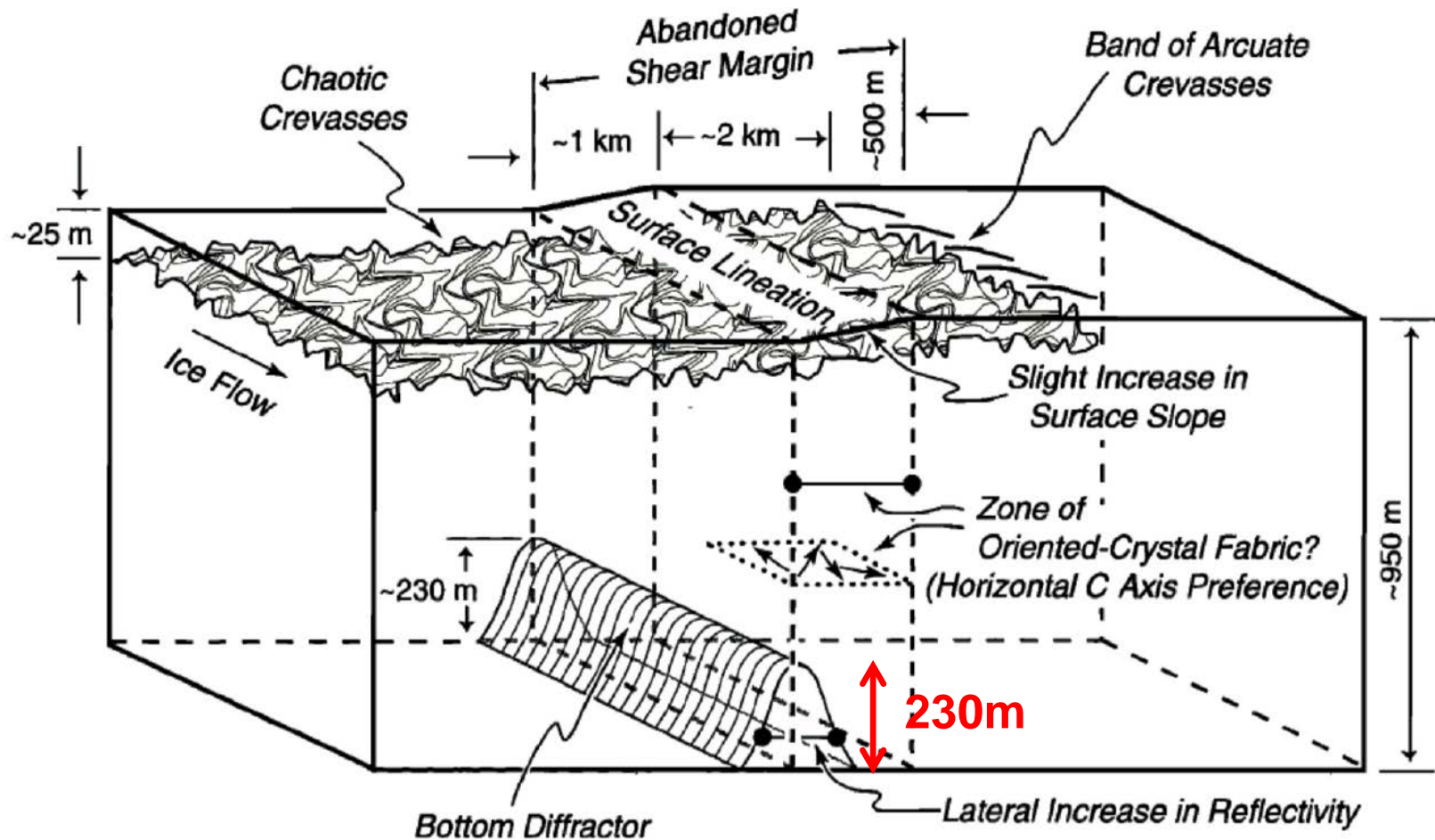


Figure by Clarke et al. 2000