

# What goes down must come up: viscoelastic deformation in the Antarctic Peninsula



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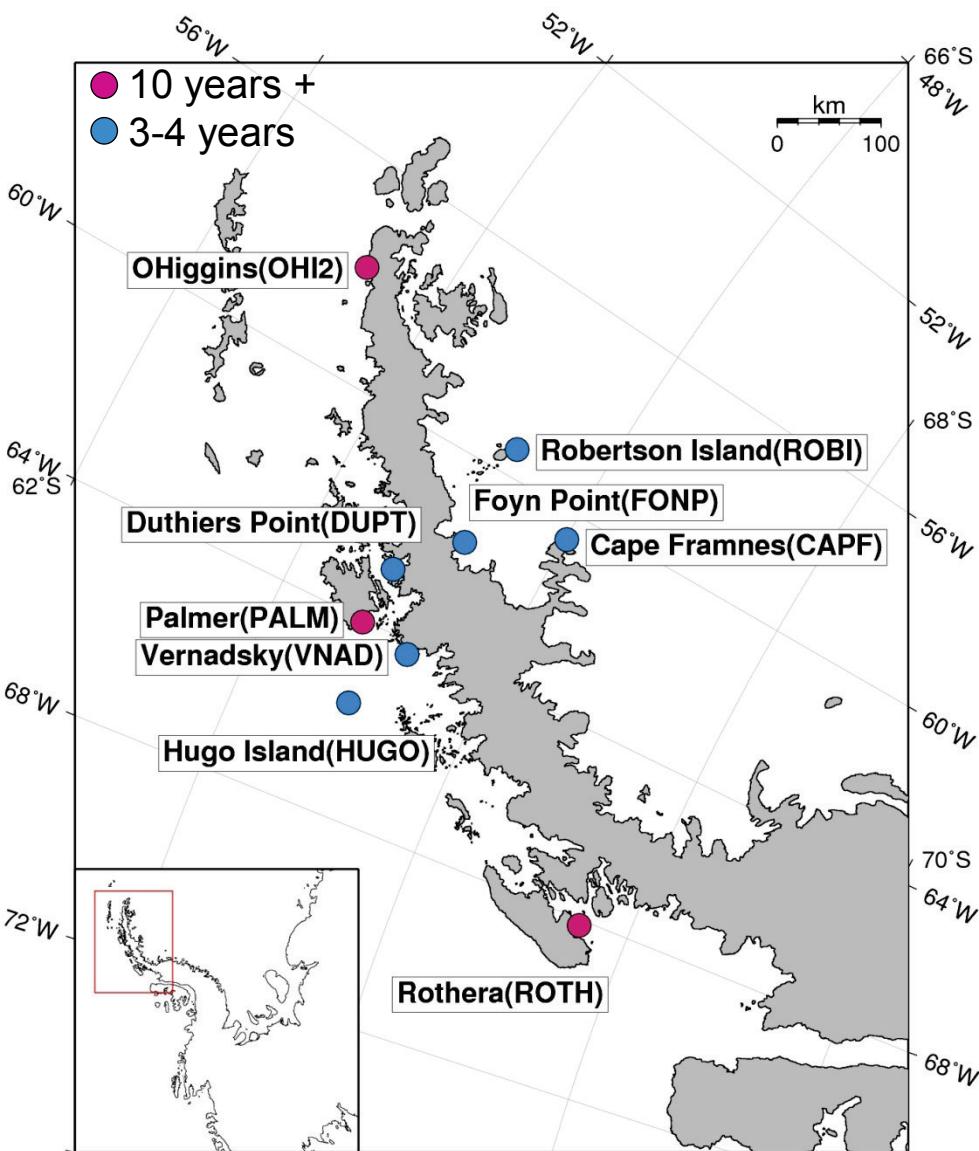
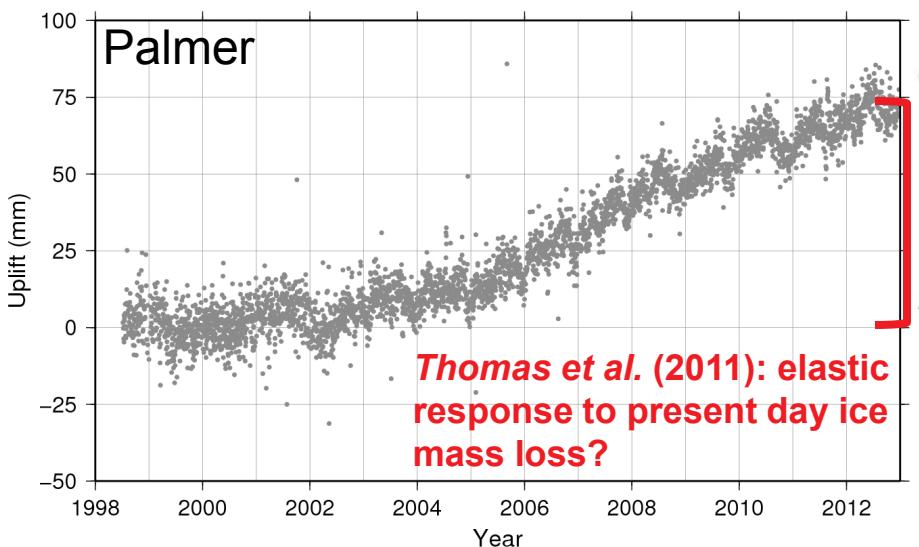
# Motivation

- How is Antarctica responding to millenial/centennial/decadal scale unloading?
- How can we explain the rapid GPS-observed uplift in the Antarctic Peninsula since ~2002?
- Well observed unloading may provide unique insight into Earth rheology (noting all GIA modelling use linear Maxwell)



# GPS Observations

- Palmer Station  
1998 – present
- LARISSA  
2009/2010 - present



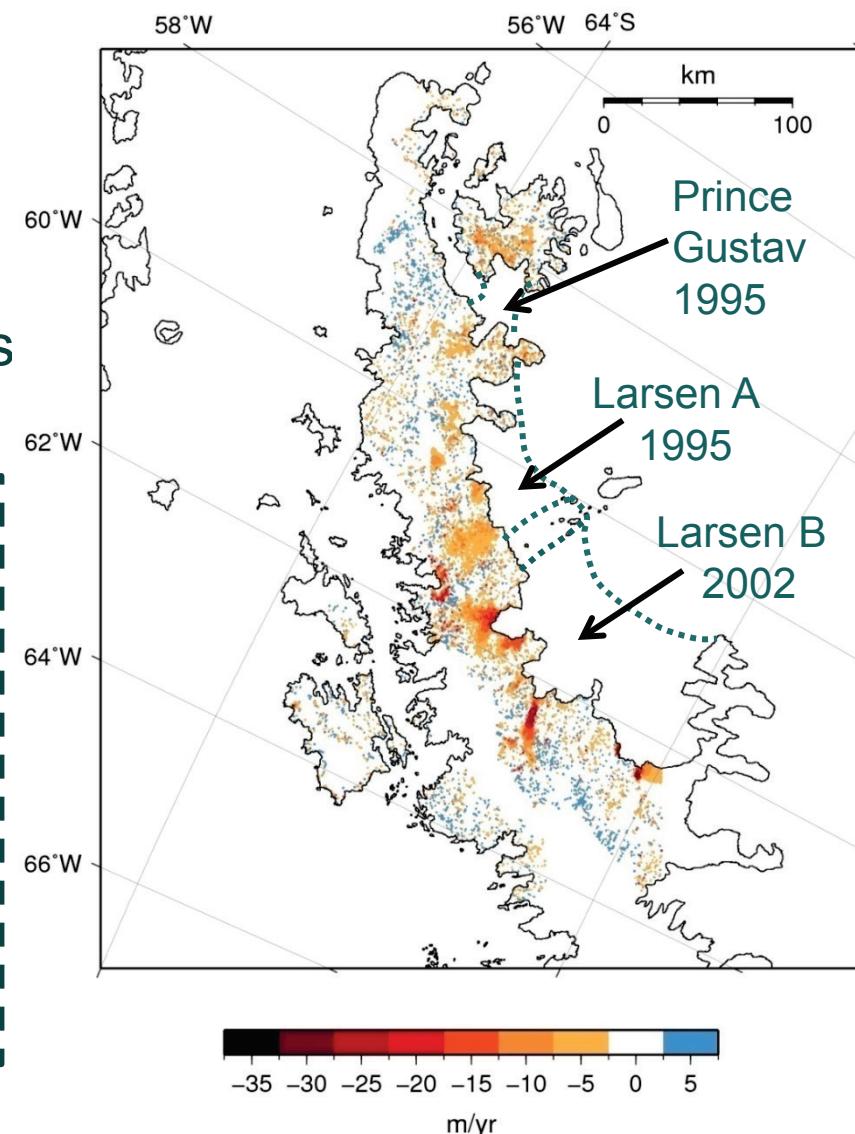
# Recent Ice Mass Changes

## Data:

- Elevation change from DEM differencing/ICESat 2003-2008
- *Berthier et al., 2012 GRL; Shuman et al., 2011 J. Glaciology; Scambos et al. (in prep)*

## Assumptions for model input:

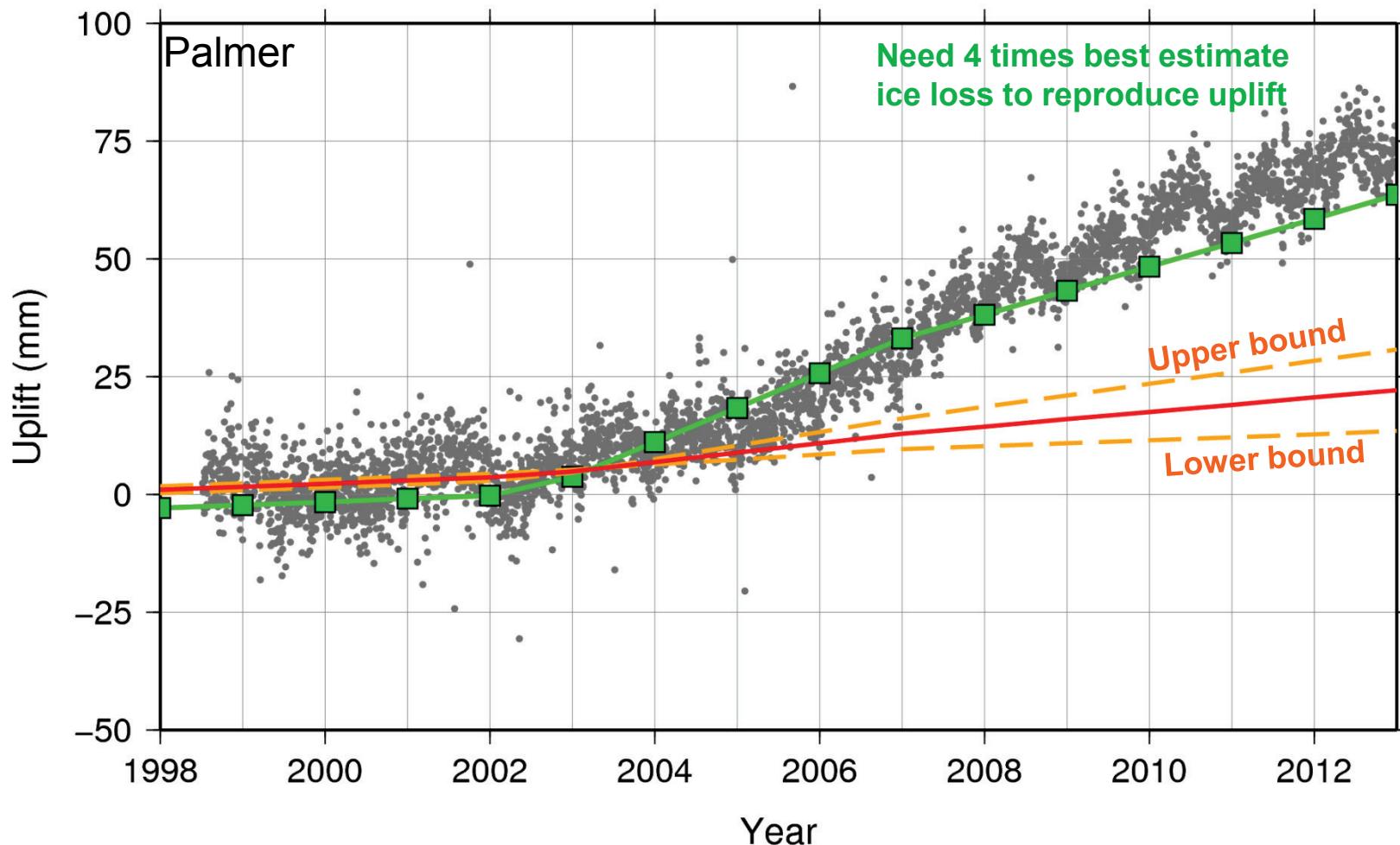
- Ice loss onset few months after collapse, negligible before this
- Mass loss continues at the same rate to present day
- Mass changes away from these glaciers is constant/long term trend



# Modelling approach

- Model solid Earth response to ice loss 1995-2013
  - Elastic response only?
  - Low viscosity mantle – viscoelastic response over ~15 years?
- Use GPS observed uplift to constrain the modelling
- Make estimate of the background rate
- Model – HresV2 - *Barletta et al.*, 2006 (GRL)

# Elastic Model - PREM

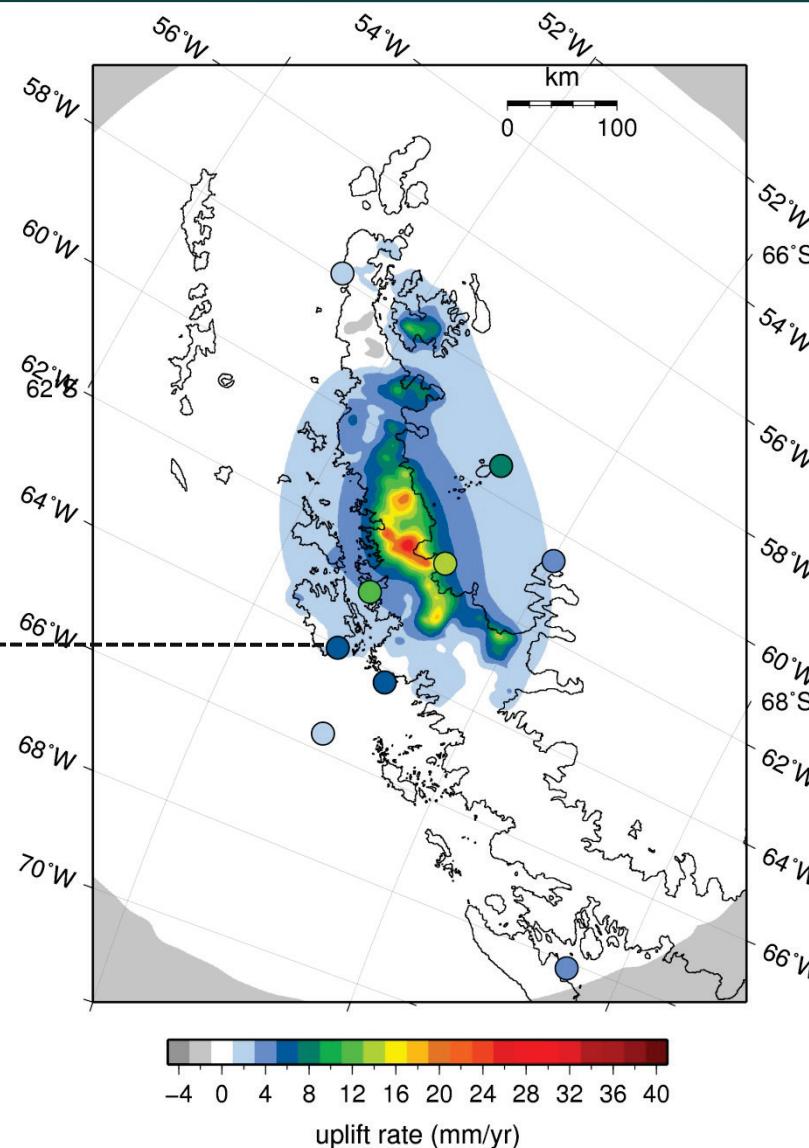


# Elastic Model – 2011 uplift rates

Modelled to  
spherical  
harmonic degree  
3754

## Palmer

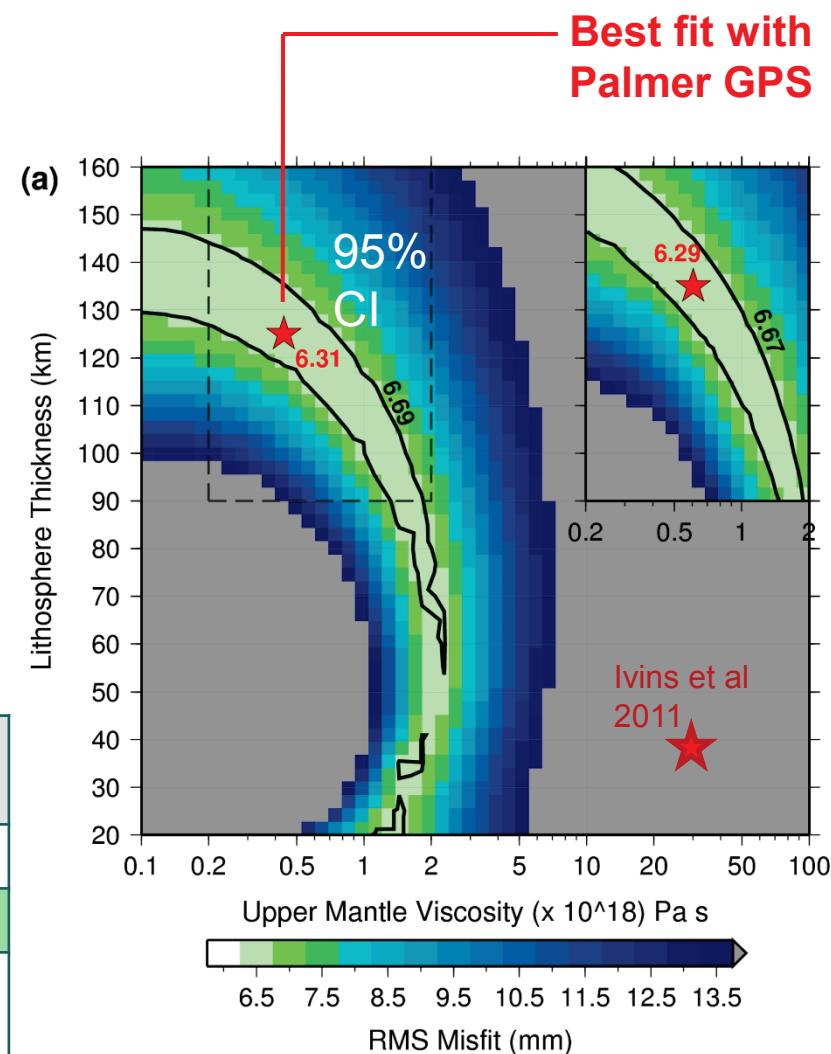
Modelled: 2.23 mm/yr  
Observed: 6.56 mm/yr



# Viscoelastic Model

- Model a viscous component
- Compute RMS misfit between modelled time series (elastic+viscous+background) and Palmer timeseries
- Incompressible Earth, with further compressible estimates shown as inset – shifts +10 to +15km

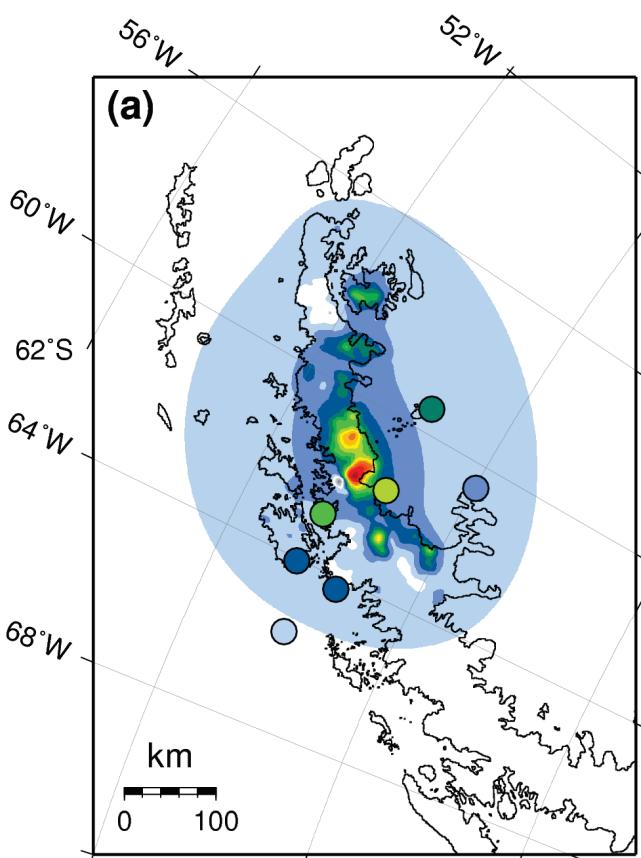
Layer	Depth to base (km)	Viscosity (Pa s)
Lithosphere	20 – 160	$1 \times 10^{51}$
Upper Mantle	400	$1 \times 10^{17} - 1 \times 10^{20}$
Transition Zone	670	$4 \times 10^{20}$ (no sensitivity)
Lower Mantle	-	$1 \times 10^{22}$ (no sensitivity)



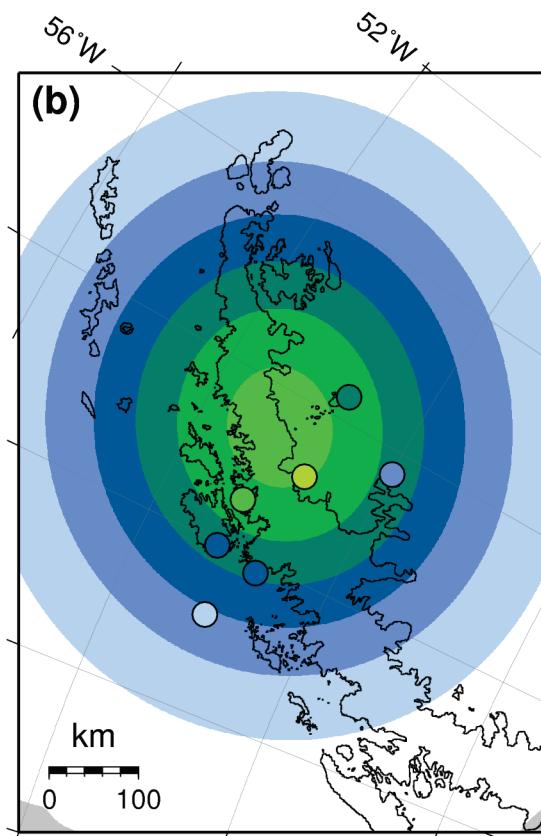
# Uplift rates for best model

Pre-1995 background rate = -2.5 mm/yr

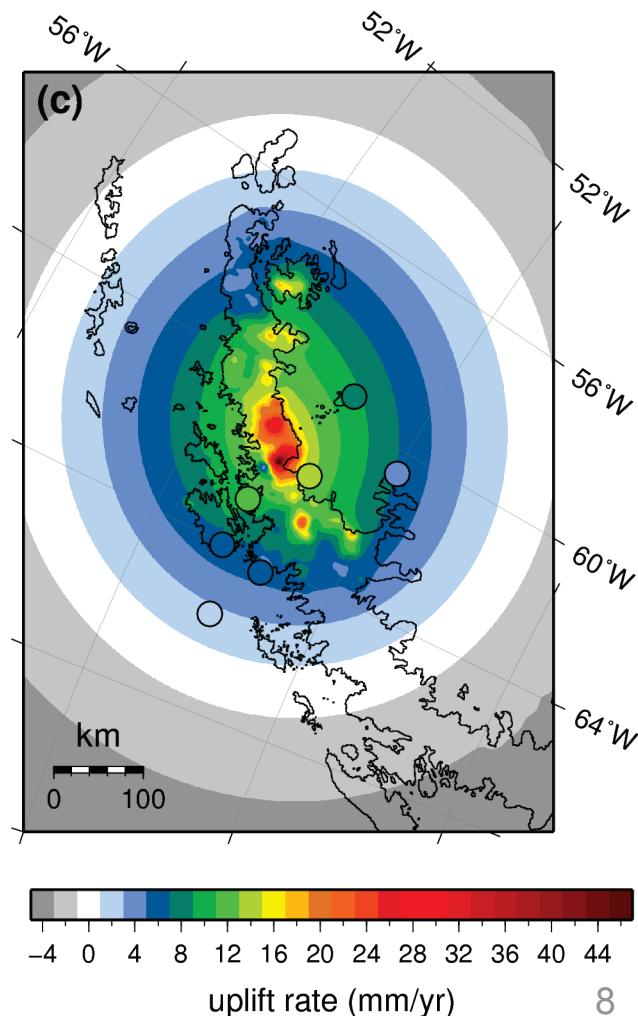
Elastic



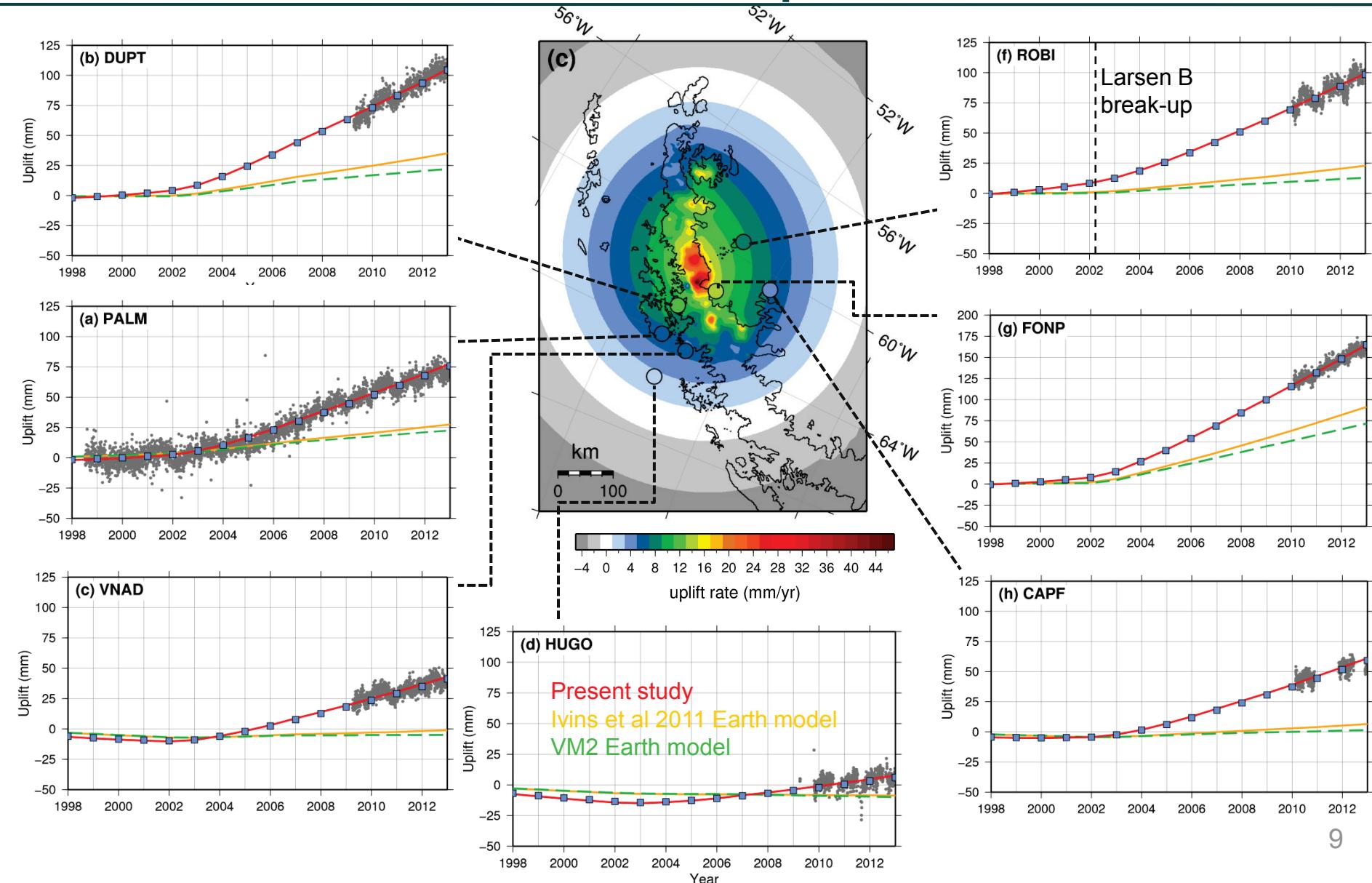
Viscous



Visco-elastic + background rate

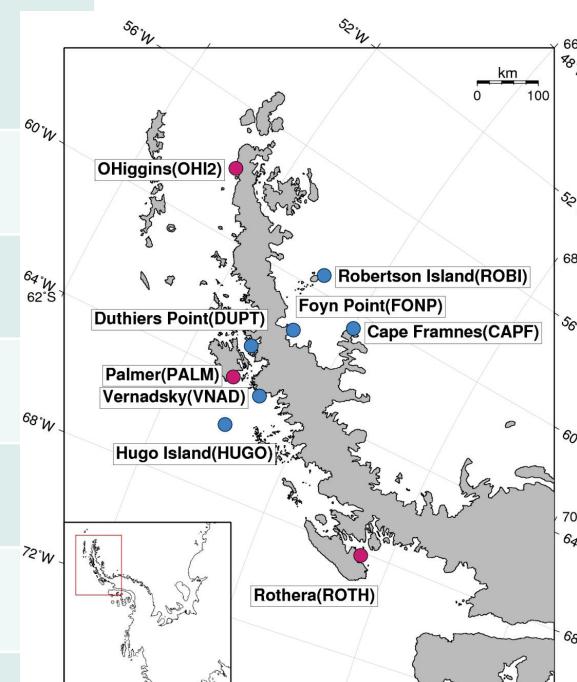


# Modelled vs. GPS uplift time series



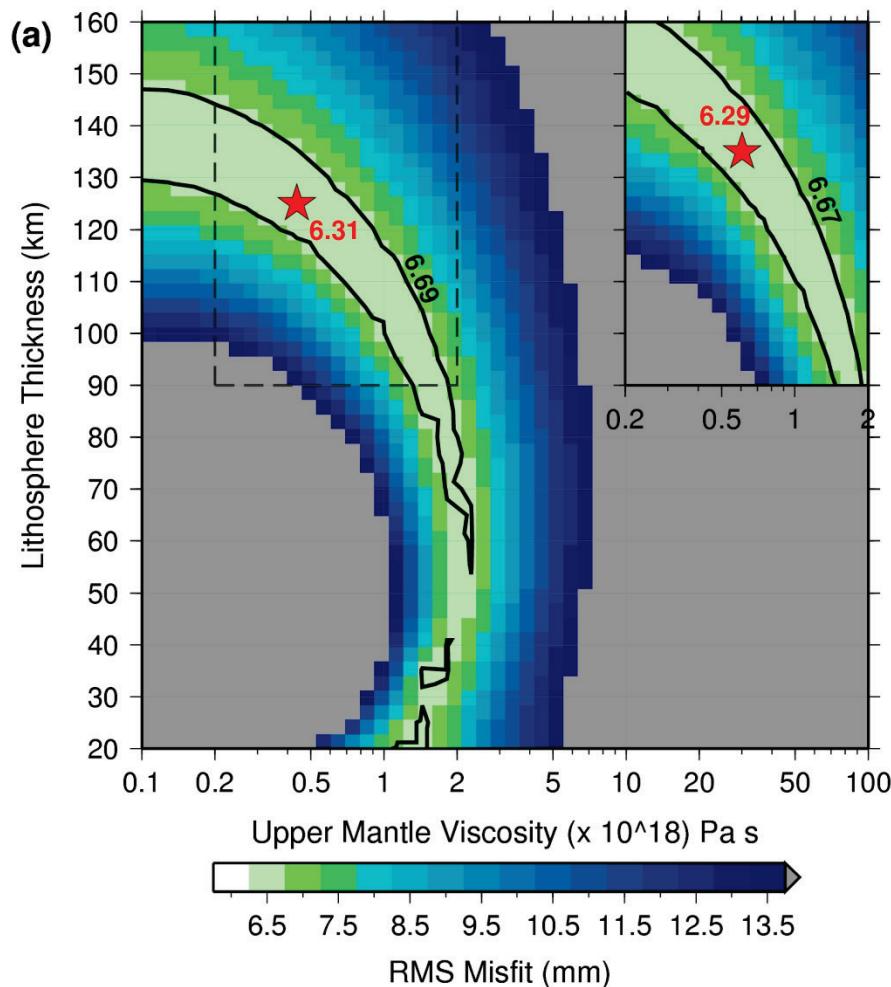
# Uplift residuals

Site	GPS Observed Uplift (mm/yr)	Viscoelastic Modelled Uplift (mm/yr)	Residual (GPS minus viscoelastic model) (mm/yr)
PALM	$6.56 \pm 1.07$ (using 2009.0 to 2013.0)	7.77	$-1.21 \pm 1.07$
CAPF	$4.50 \pm 1.45$	7.09	$-2.60 \pm 1.45$
DUPT	$12.84 \pm 1.06$	10.08	$2.76 \pm 1.06$
FONP	$14.89 \pm 1.35$	15.98	$-1.08 \pm 1.35$
HUGO	$1.69 \pm 1.66$	2.82	$-1.13 \pm 1.66$
ROBI	$7.79 \pm 1.45$	9.52	$-1.73 \pm 1.45$
VNAD	$5.78 \pm 1.21$	5.77	$0.01 \pm 1.21$

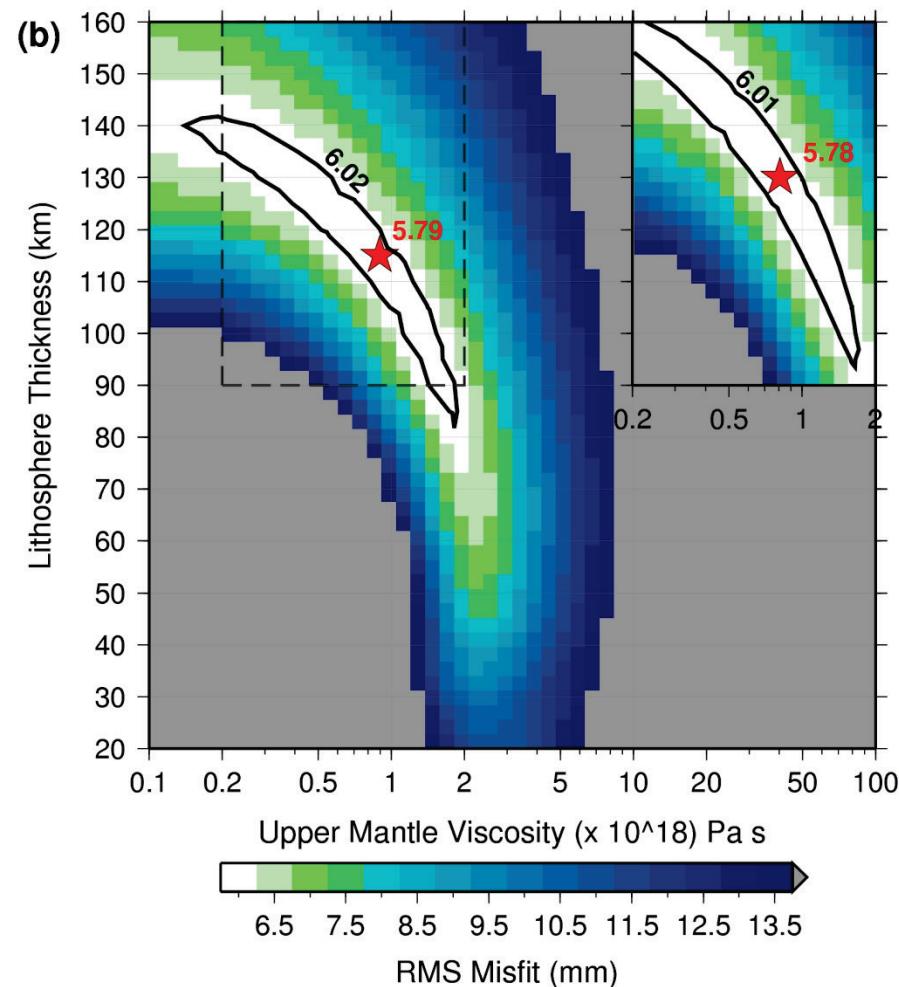


# Viscoelastic Model

Best fit with  
Palmer GPS

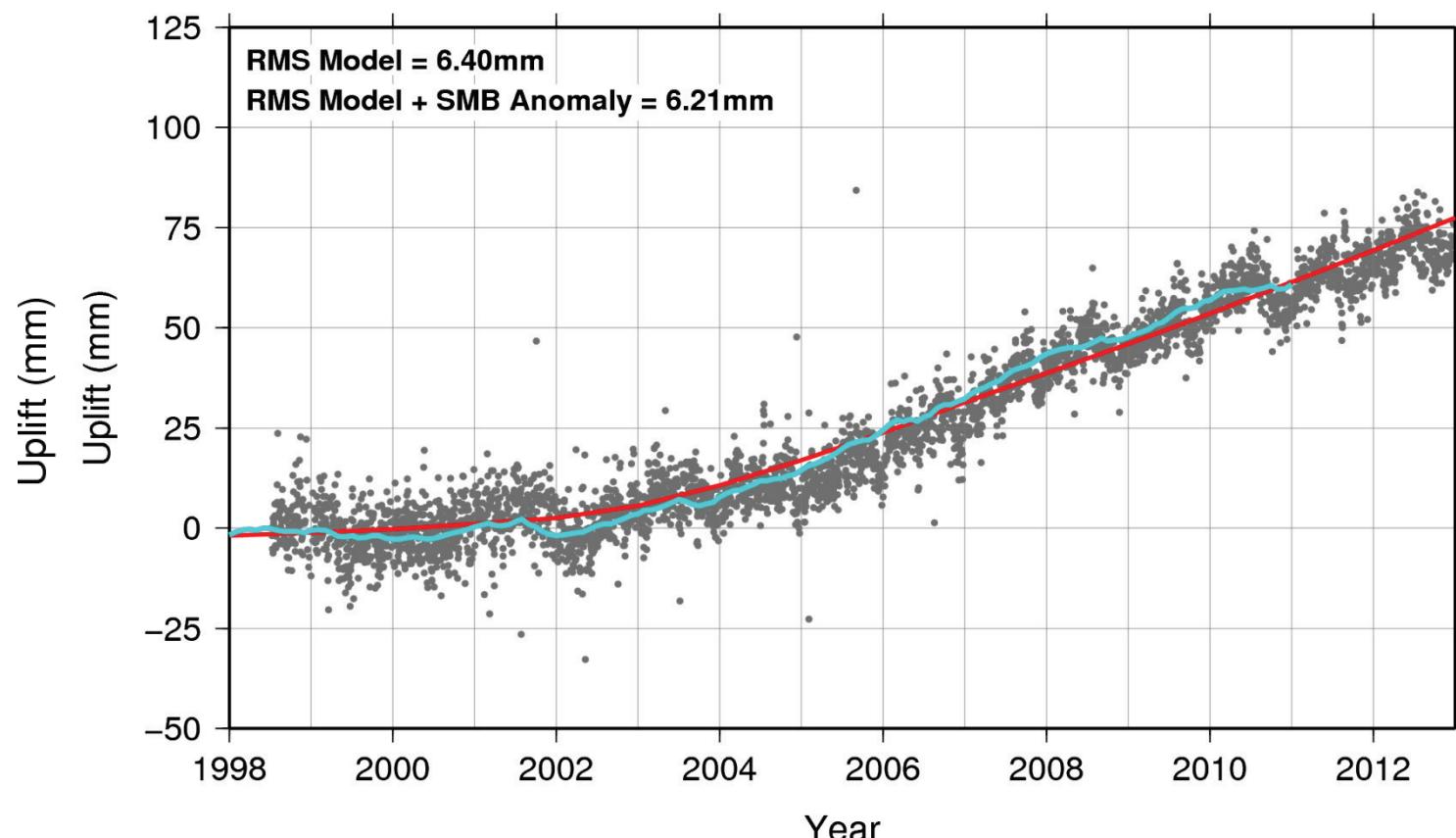


Best fit with  
all GPS



# Local time-variable elastic effects?

- SMB anomalies may cause an additional elastic effect.
- Test using RACMO2.1/ANT (Lenaerts et al., 2012) SMB, anomaly to 1979 - 2011.0 mean

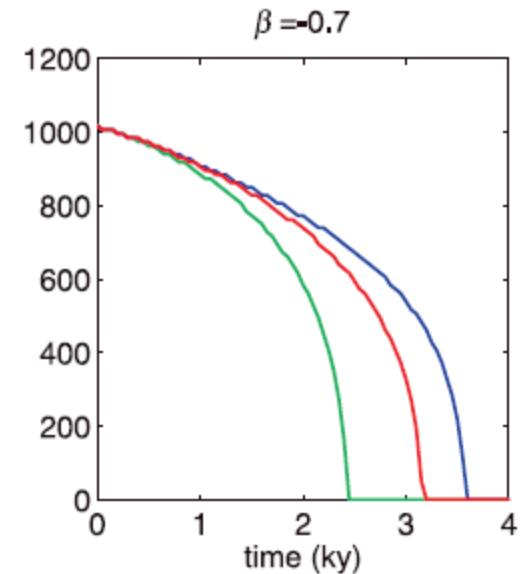
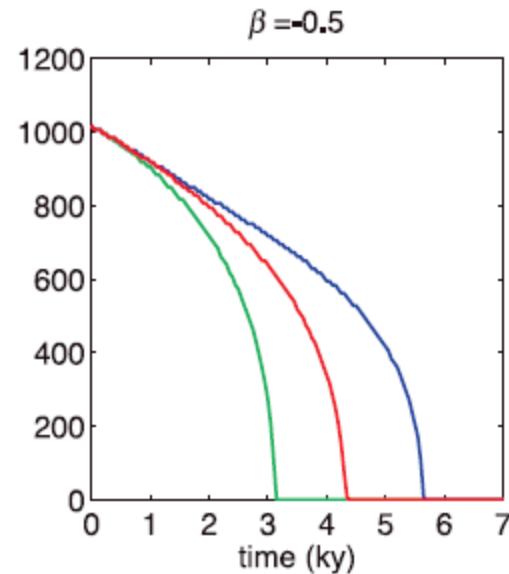
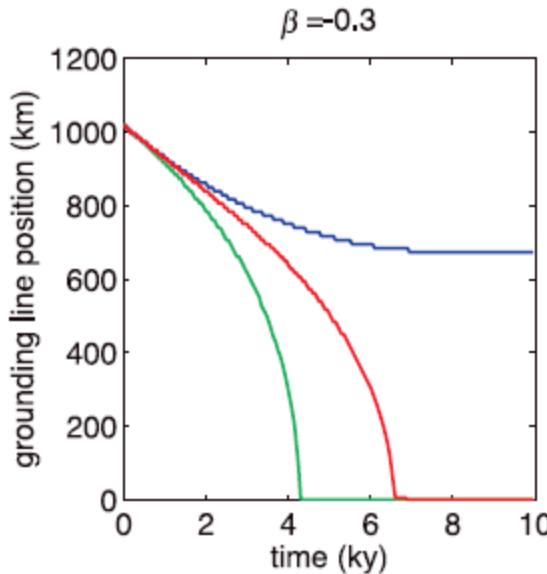


# Summary

- A unique experiment – needs to be ongoing [dh/dt and new GPS to capture before/after future events e.g., SCAR inlet]
- Elastic response alone cannot explain the observed uplift & viscous component is required
- Modelling suggests the range of parameters that fit the GPS observations are:
  - Lithospheric thickness: ~70-135km
  - Upper mantle viscosity:  $< 2 \times 10^{18}$  Pa s
- Pre-2002 uplift at PALM very low. Pre-1995 uplift likely negative [-2.5mm/yr]
- Earth models of Peltier (VM2) and Ivins et al 2011 do not provide good fits to the GPS when combined with post-1995 unloading – is time-dependent rheology at play?

# Rapid viscous uplift and WAIS

- Gomez et al 2012 highlighted importance of viscoelastic (blue) response to reverse slope beds ( $\beta$ = slope m/km) cf not considering (green)
- Elastic component (red) dominates for first several centuries

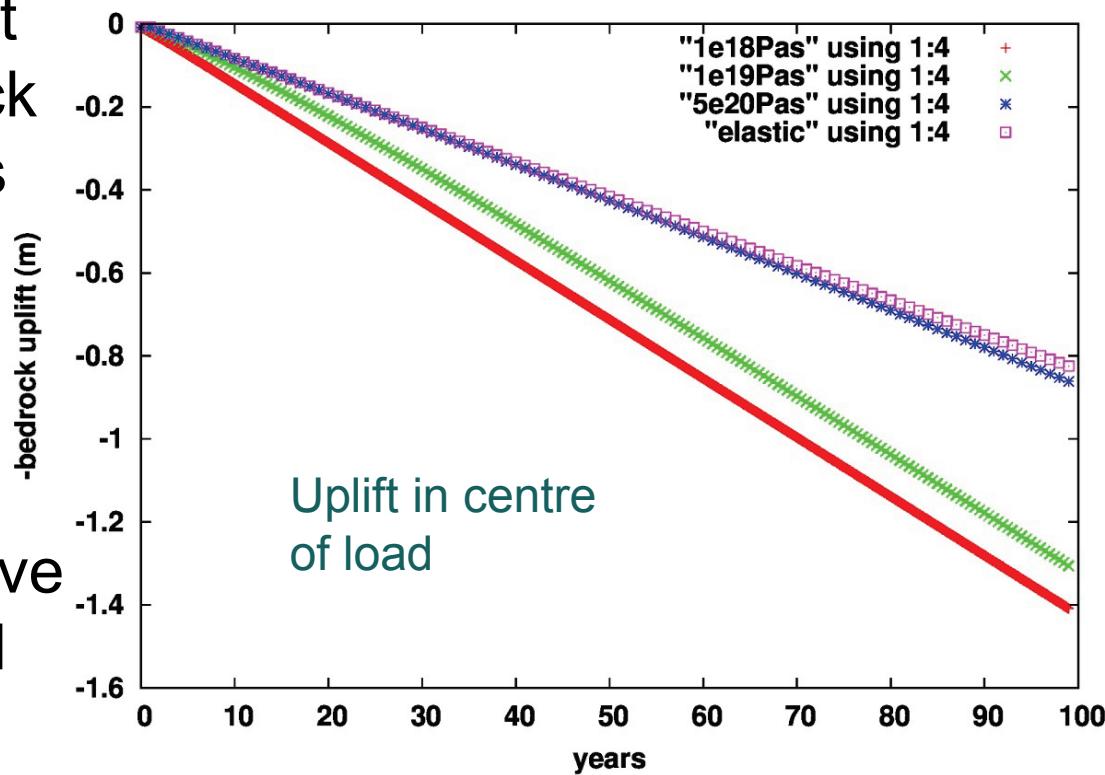


# Rapid viscous response and grounding line stabilisation

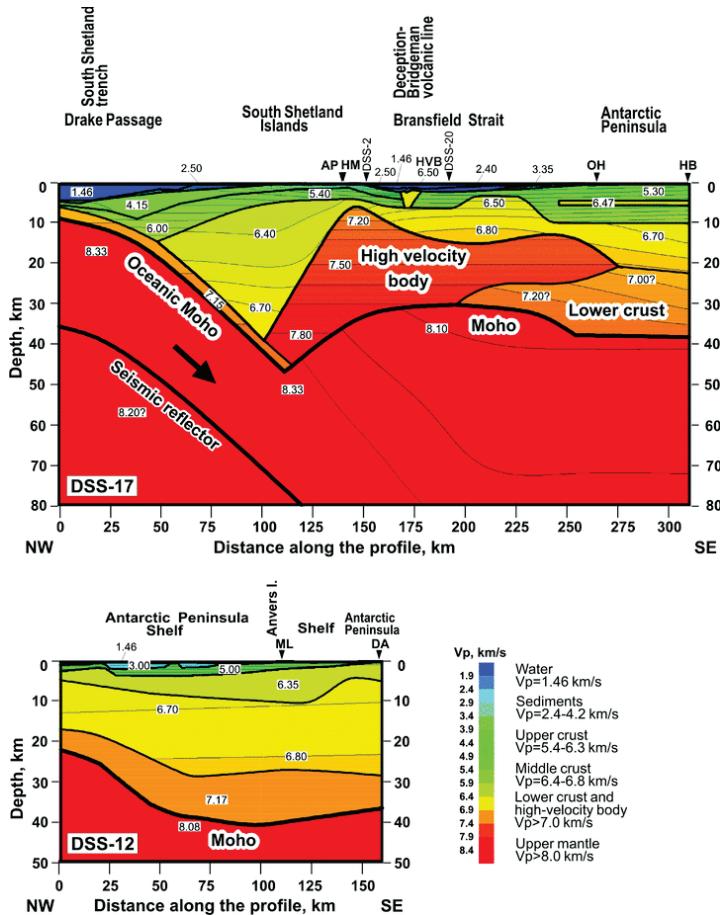
- But Gomez et al adopted globally typical  $\eta_{UM}=5\times10^{20}\text{Pa s}$  ( $t_m=236\text{y}$ ) and elastic lithosphere 120km
- Weaker Earth model may apply in WAIS, possibly locally very weak
- Examine PIG-like scenario – 80x20km region thinning at 2m/yr from t=0 & ongoing
- Semi-infinite half-space with elastic lid of 50km and various  $\eta$  and linear and power-law within RELAX

# Rapid viscous response and grounding line stabilisation

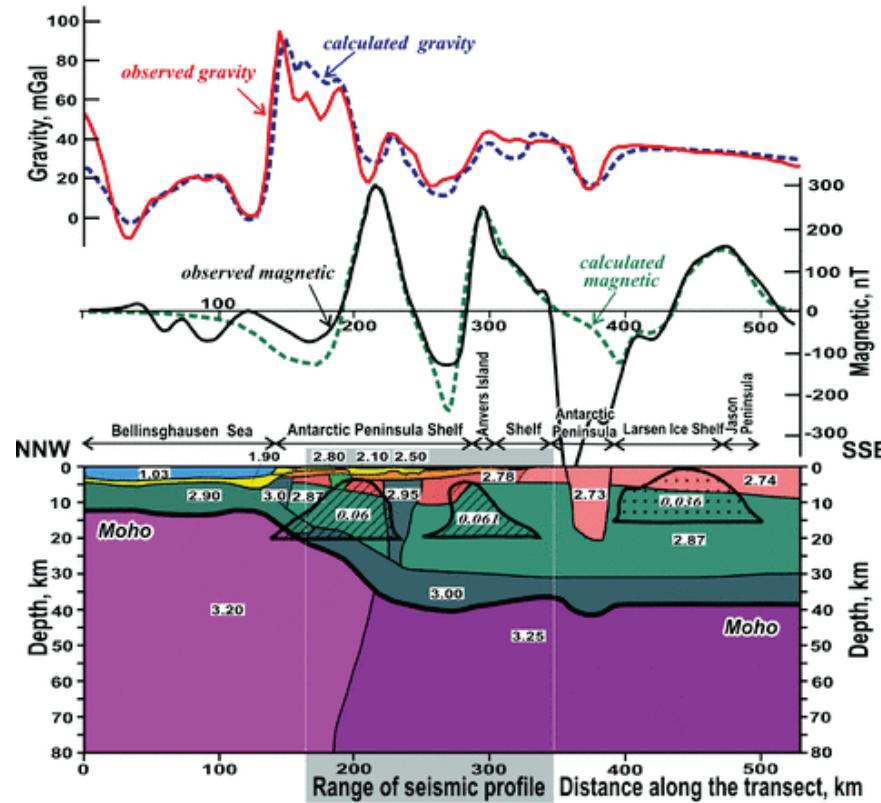
- Weak Earth could result in ~50% greater bedrock adjustment within years to decades cf elastic or high viscosity model
- not shown: Power-law rheology concentrates uplift temporally (effective viscosity reduction) and spatially near to load change so important for bedrock slopes



# Joint geophysical and petrological models for the lithosphere structure of the Antarctic Peninsula continental margin



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